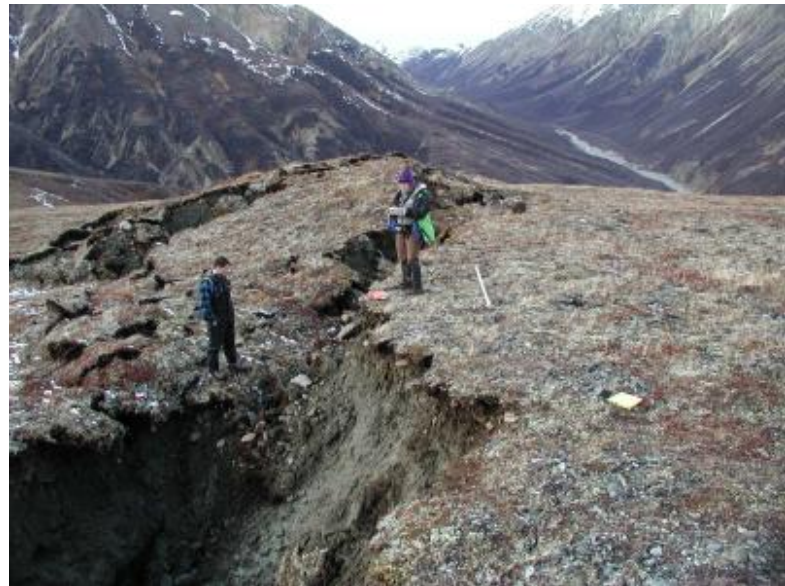
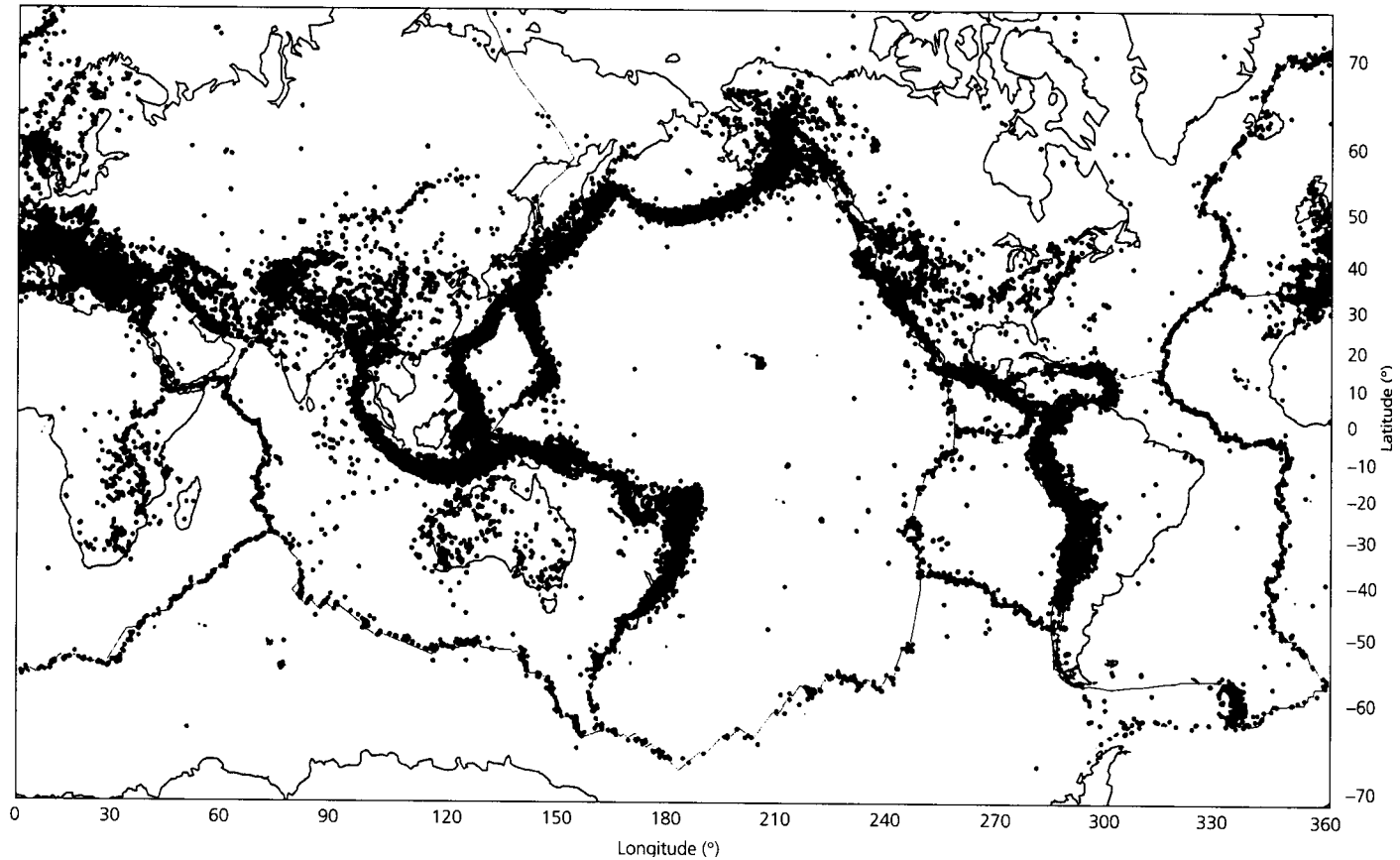


Lecture 8 – Seismo-tectonics

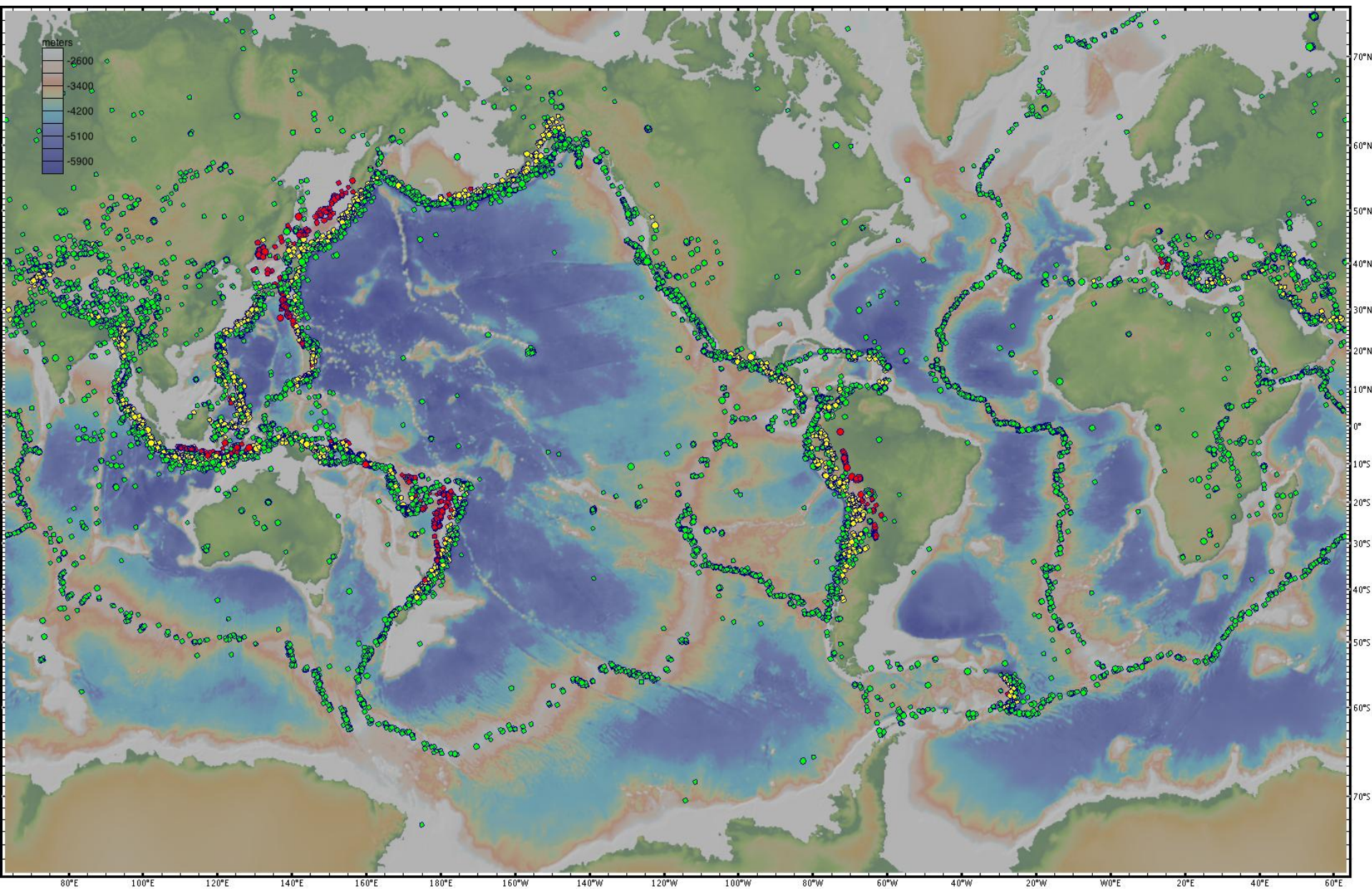
read KKV ch. 5.2

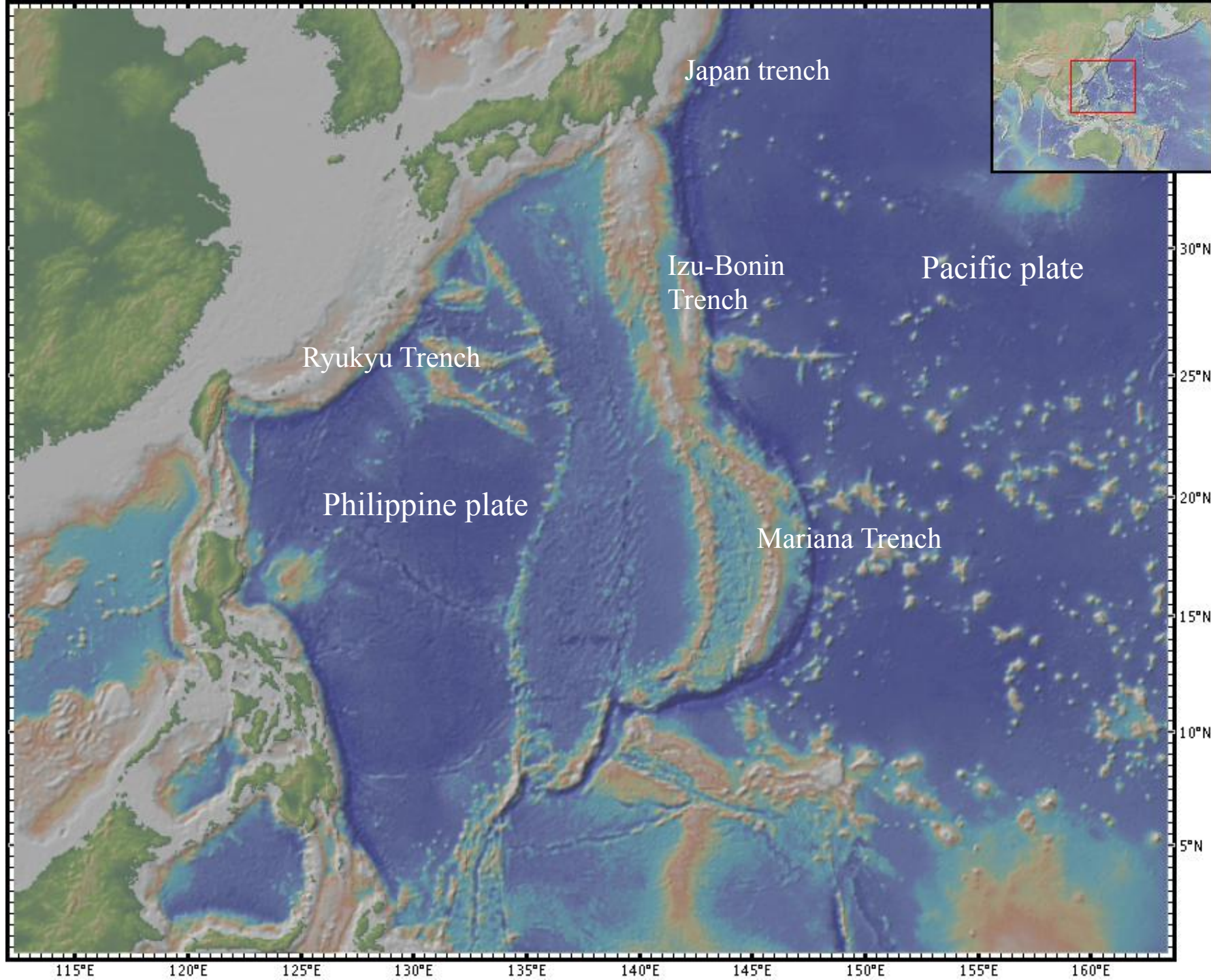


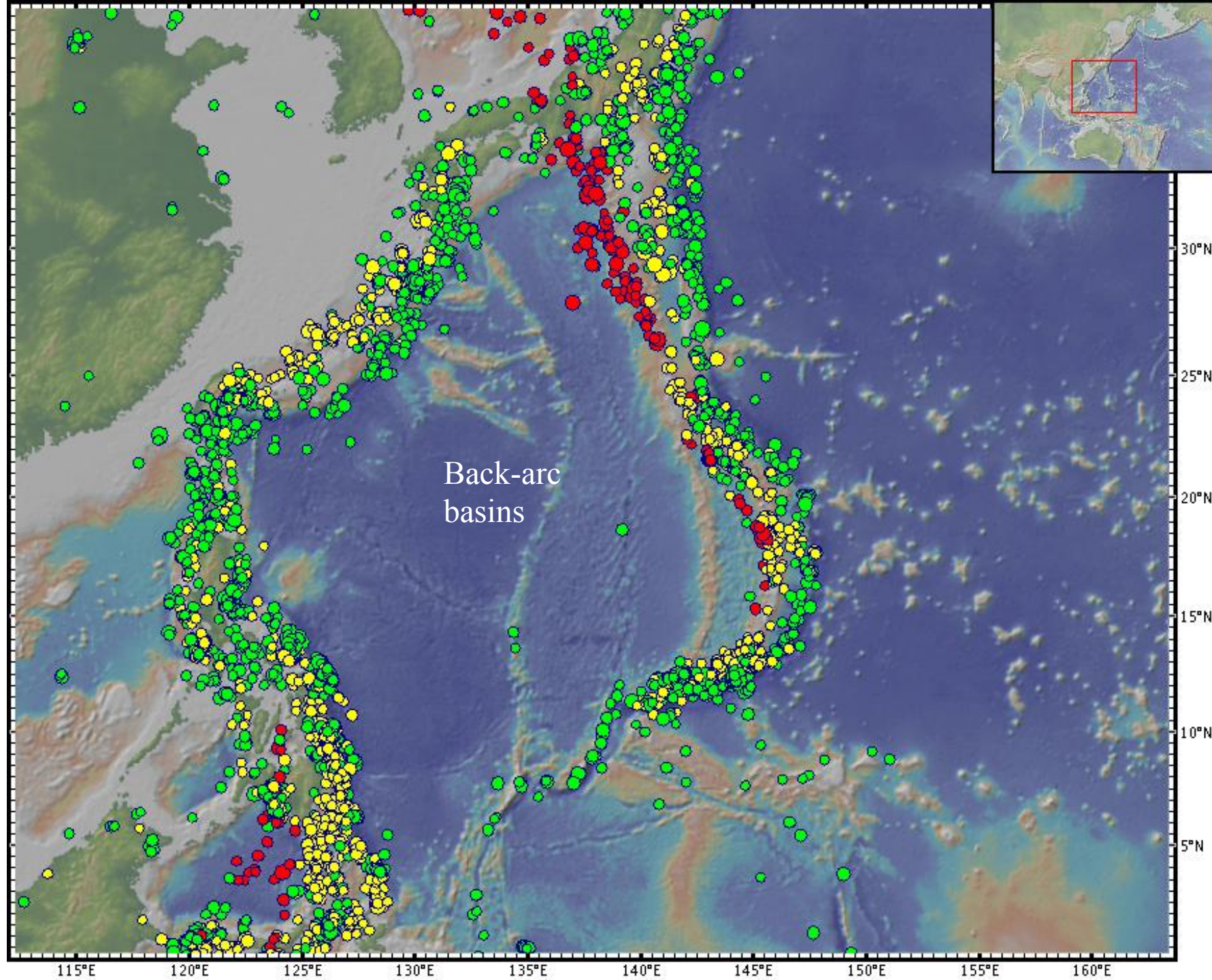
Global distribution of earthquakes



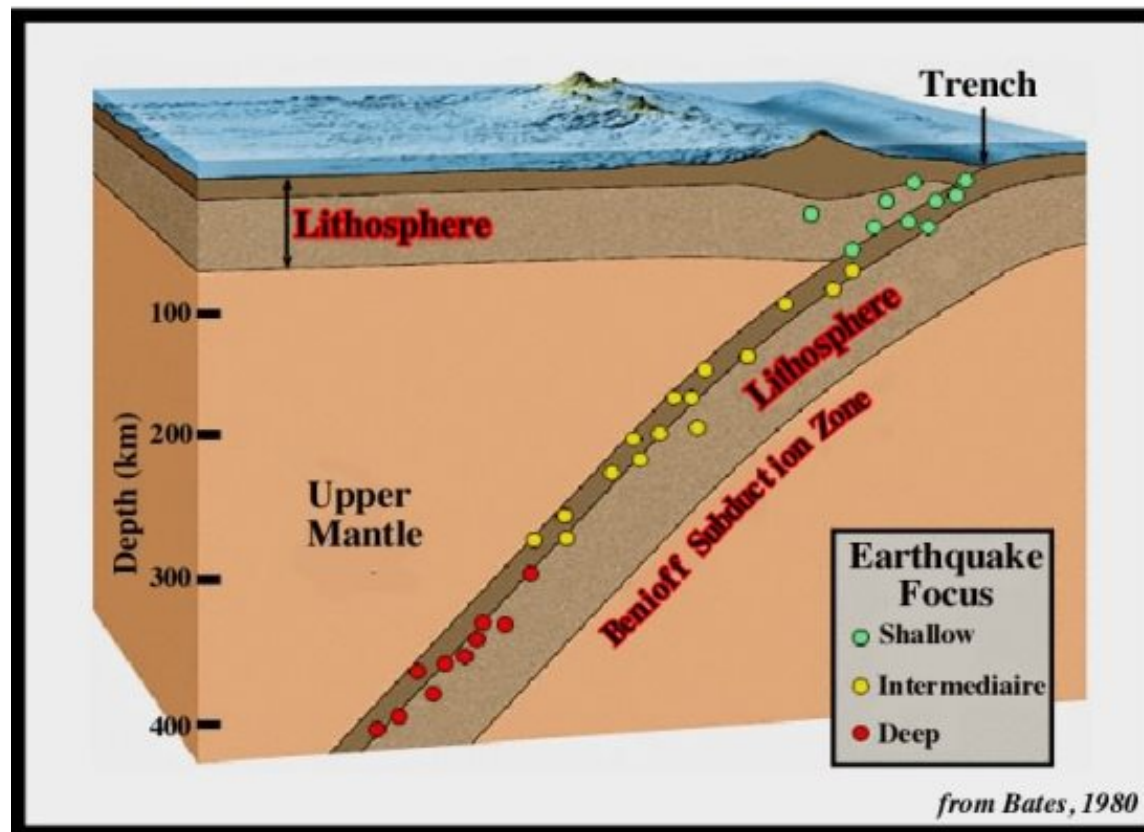
Focal Depths: green < 50 km, yellow = 50 to 250 km, red > 250 km



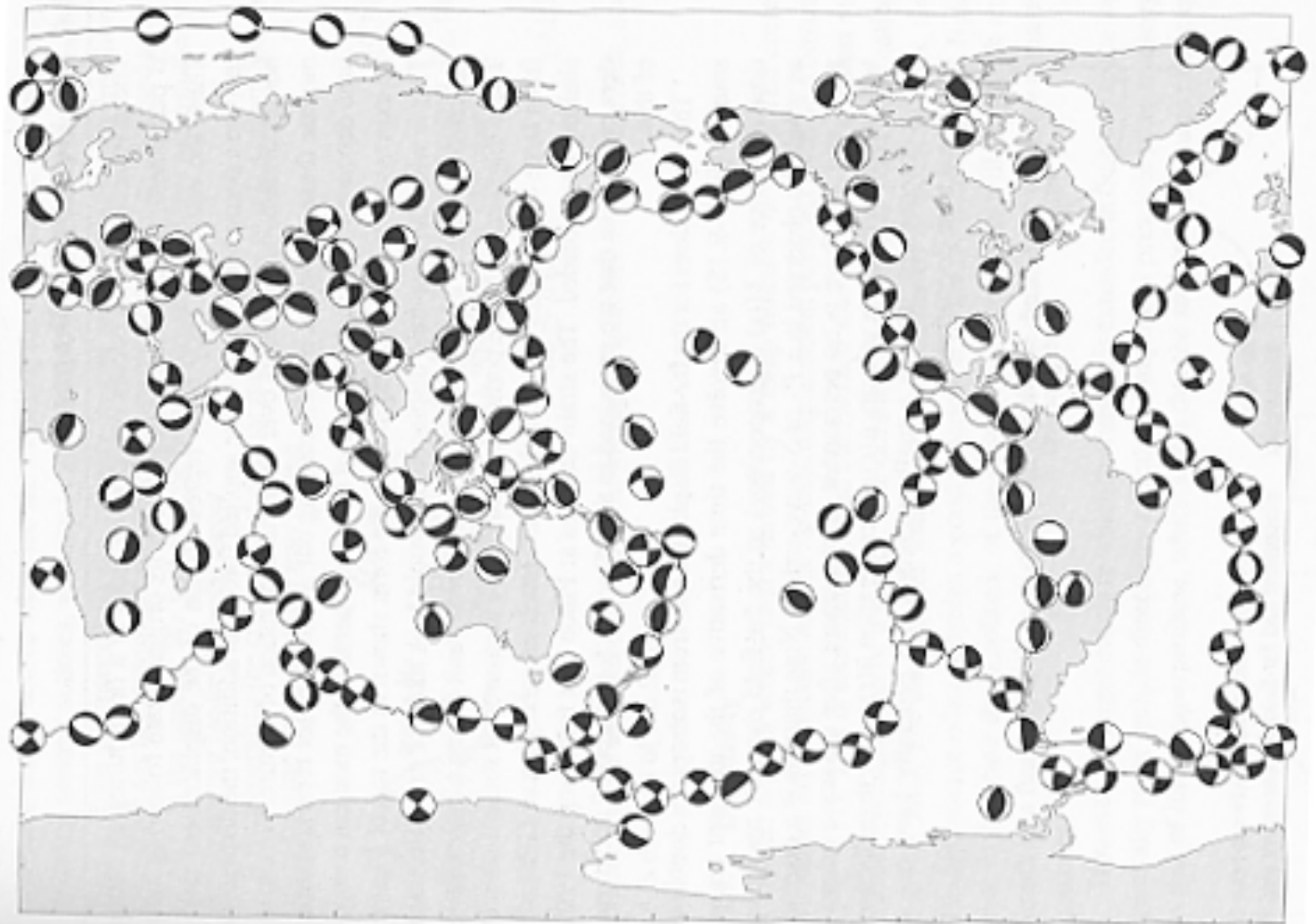


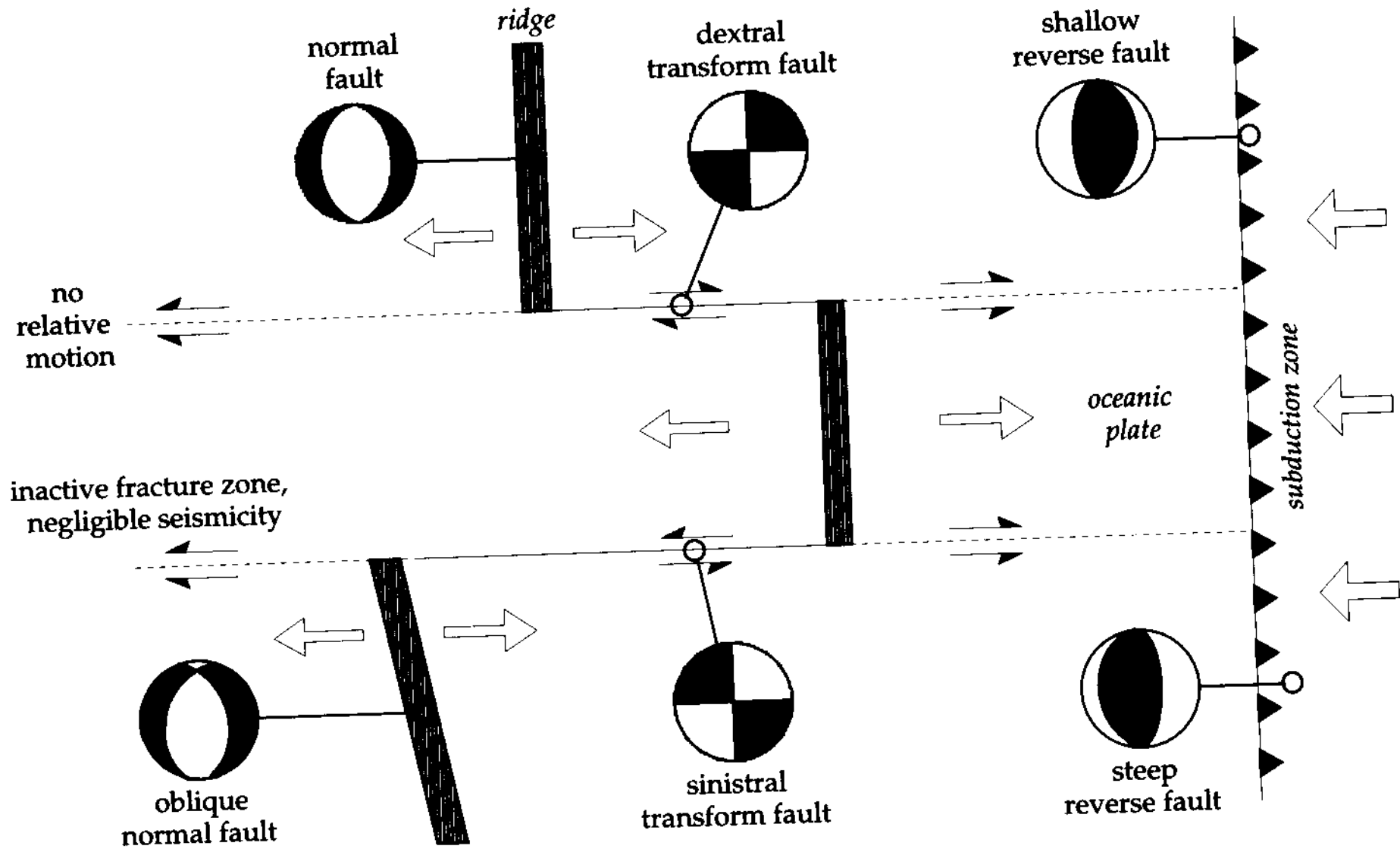


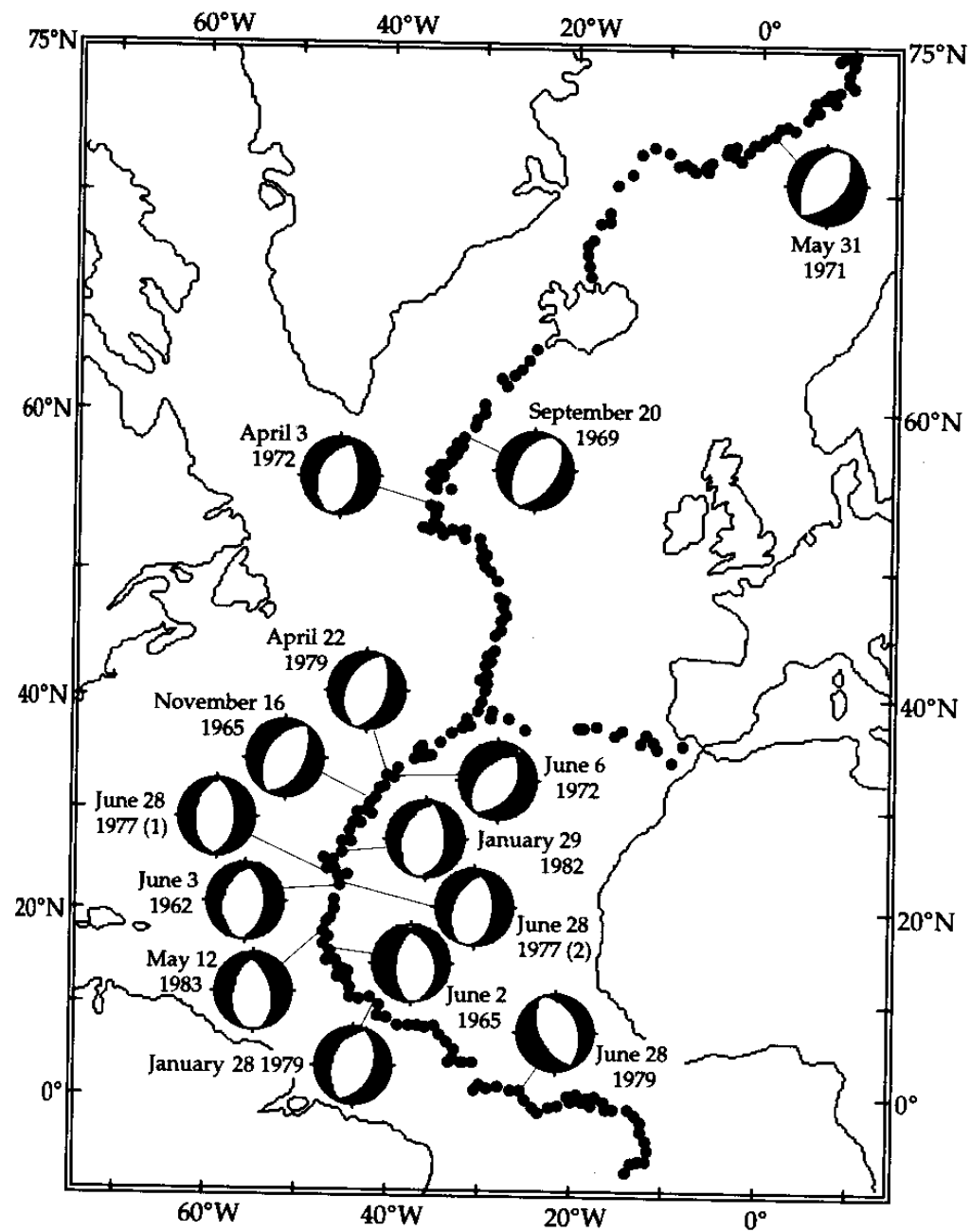
Benioff zone

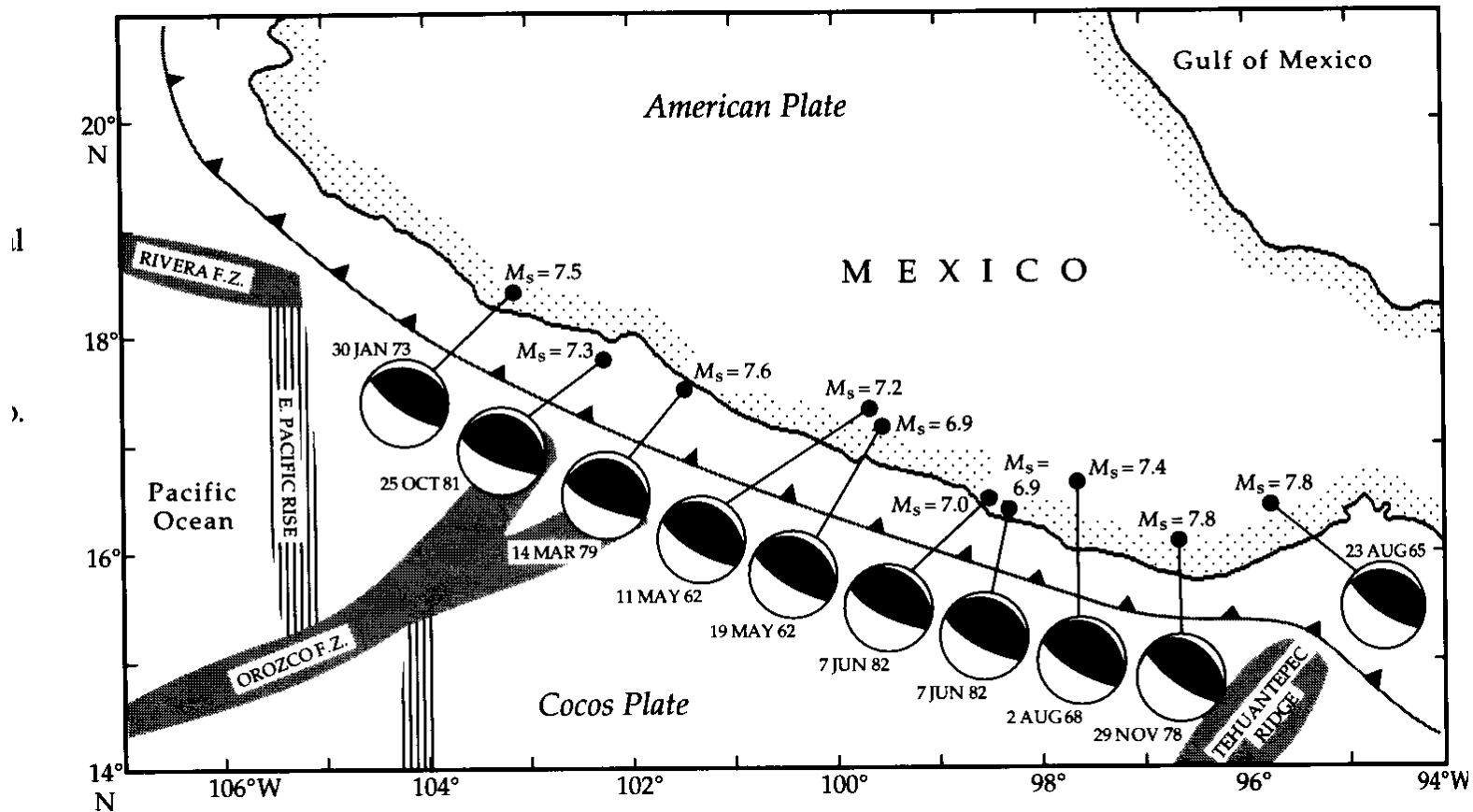


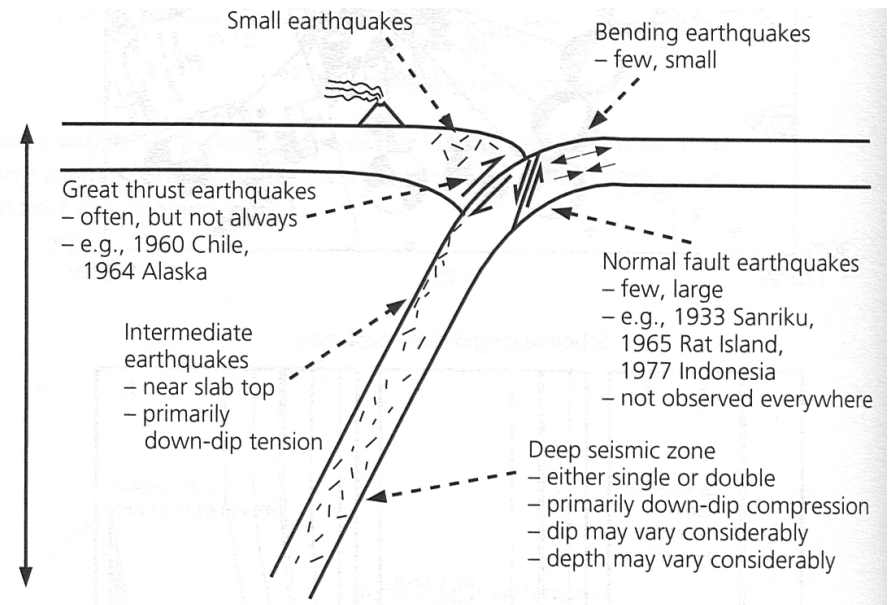
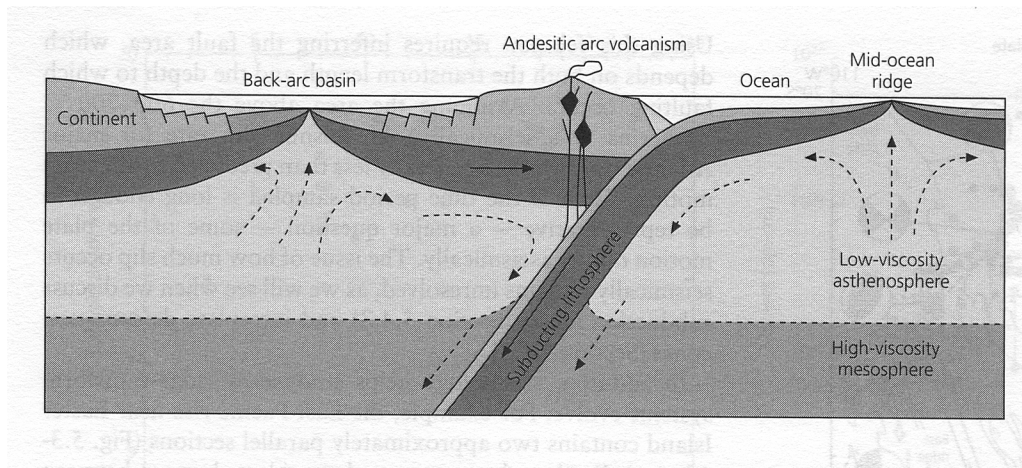
Global distribution of the earthquake focal mechanisms



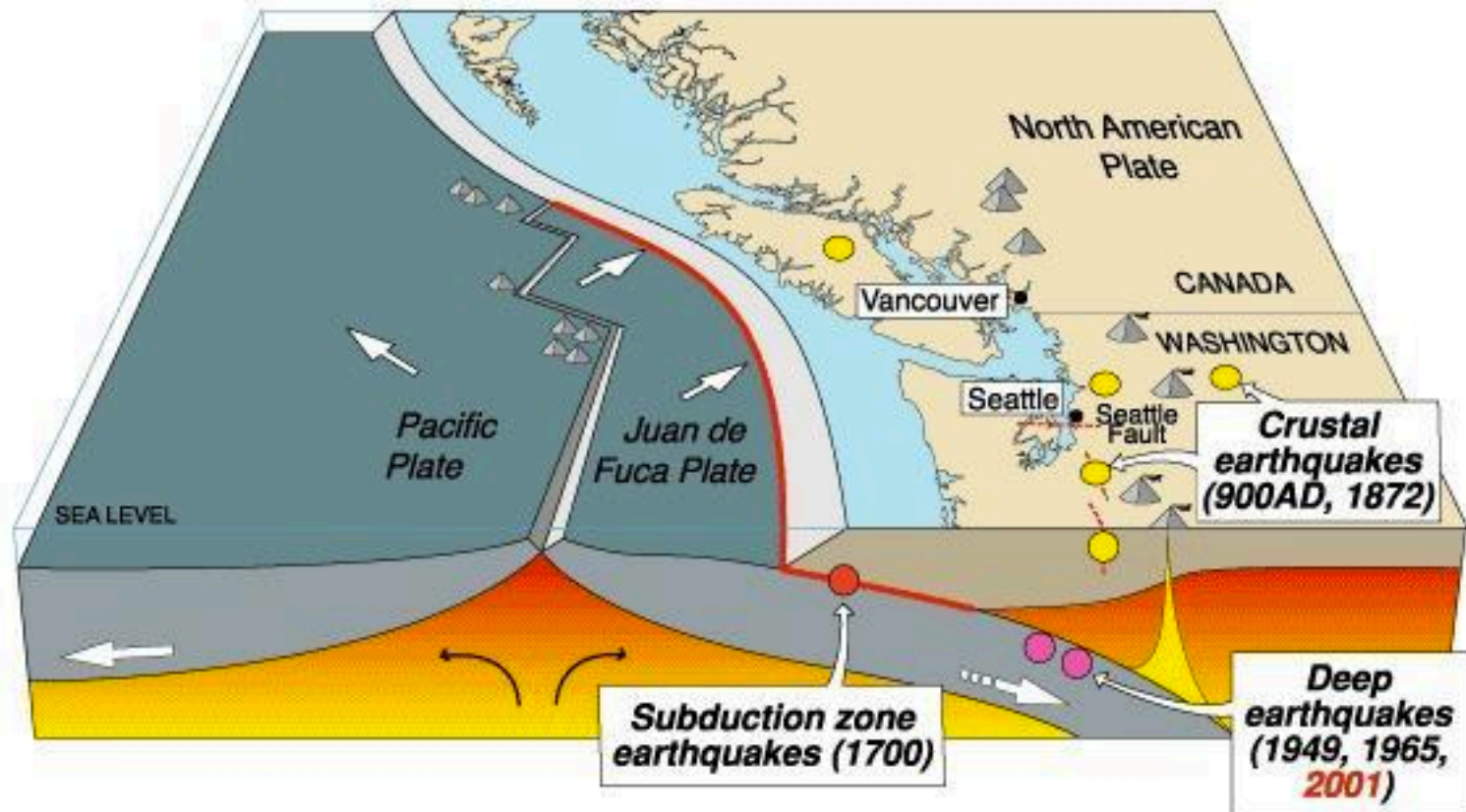








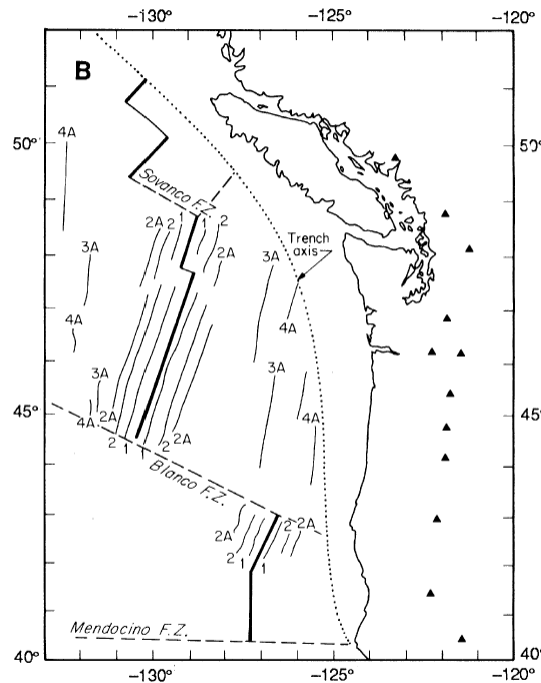
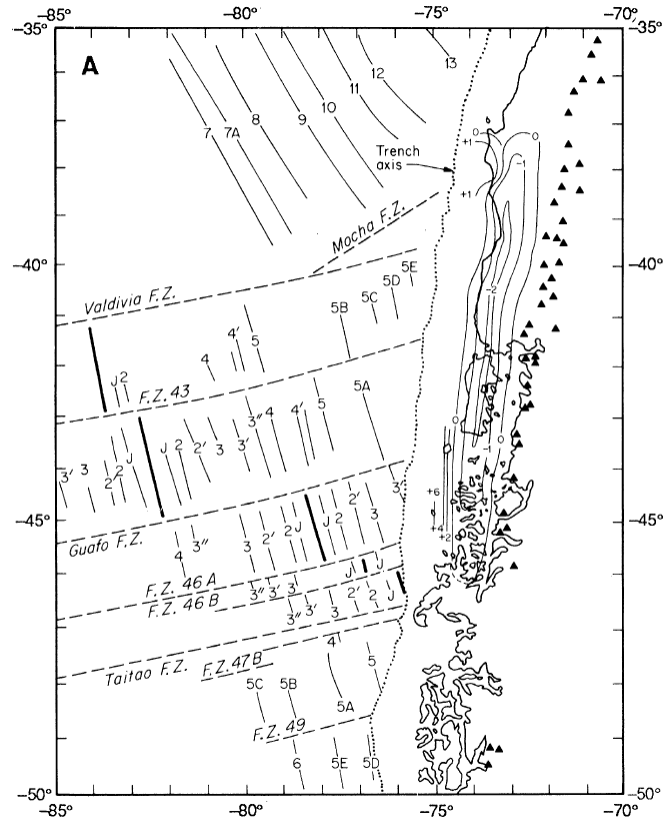
Cascadia earthquake sources



Earthquake Hazards on the Cascadia Subduction Zone

THOMAS H. HEATON AND STEPHEN H. HARTZELL (1987)

Large subduction earthquakes on the Cascadia subduction zone pose a potential seismic hazard. Very young oceanic lithosphere (10 million years old) is being subducted beneath North America at a rate of approximately 4 centimeters per year. The Cascadia subduction zone shares many characteristics with subduction zones in southern Chile, southwestern Japan, and Colombia,



Similarities between Southern Chile trench and Cascadia:

Is subduction of young crust a predictor of large earthquakes?

(But, 120 Ma crust involved in Japan 2011)

THE REALLY BIG ONE

An earthquake will destroy a sizable portion of the coastal Northwest. The question is when.

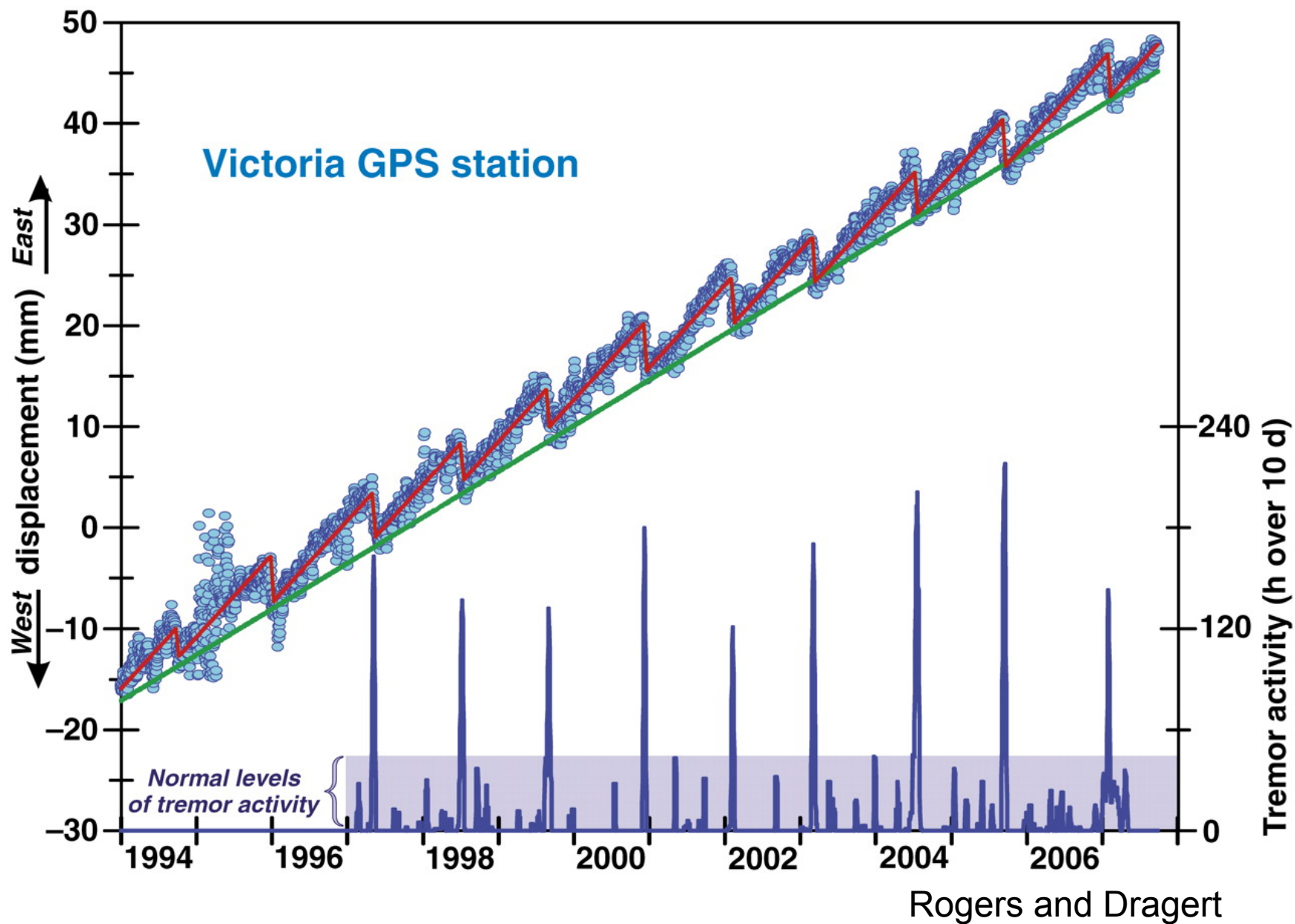


By Kathryn Schulz

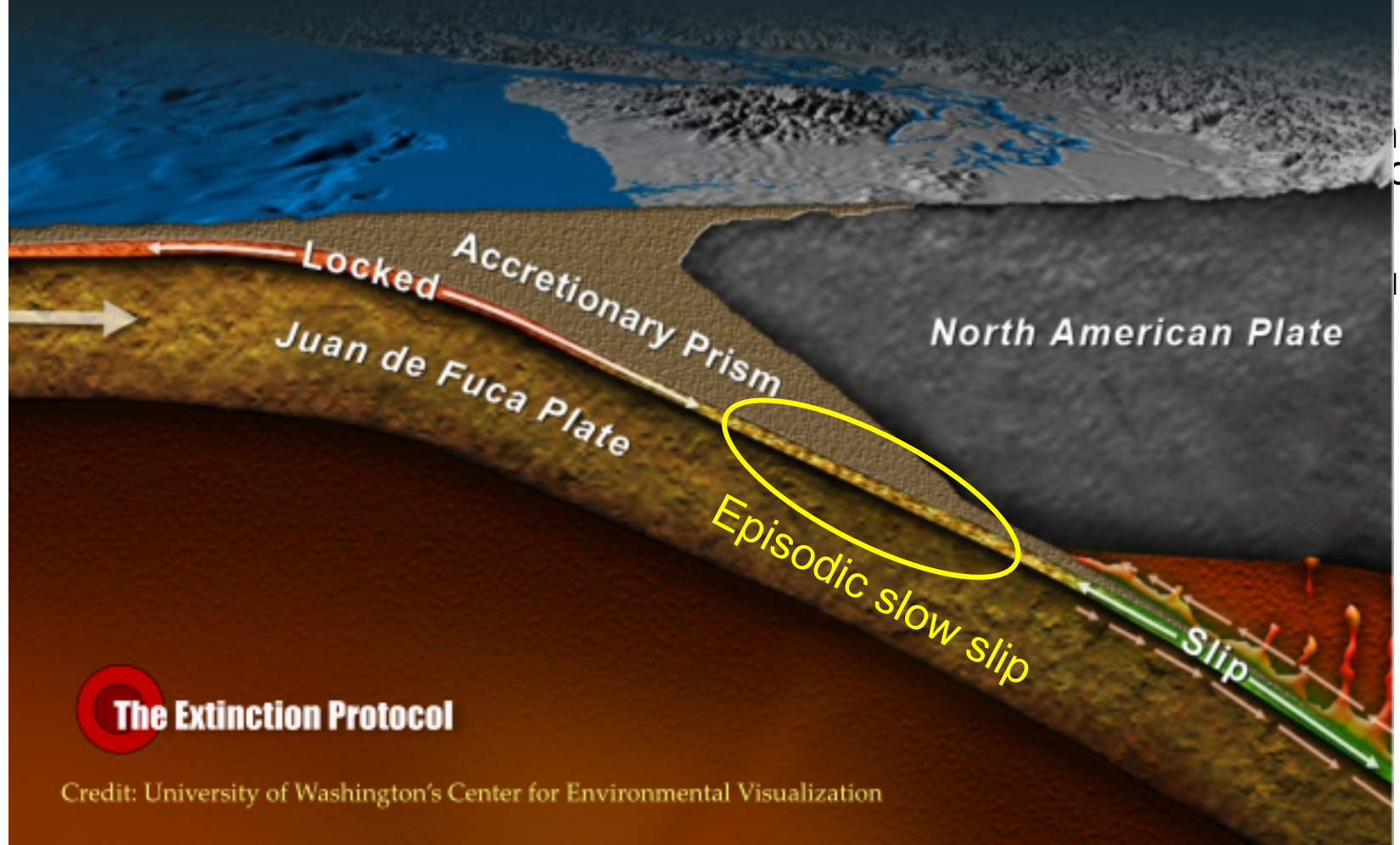
When the 2011 earthquake and tsunami struck Tohoku, Japan, Chris Goldfinger was two hundred miles away, in the city of Kashiwa, at an international meeting on seismology. As the shaking started, everyone in the room began to laugh. Earthquakes are common in Japan—that one was the third of the week—and the participants were, after all, at a seismology conference. Then everyone in the room checked the time.



Slow slip and tremor



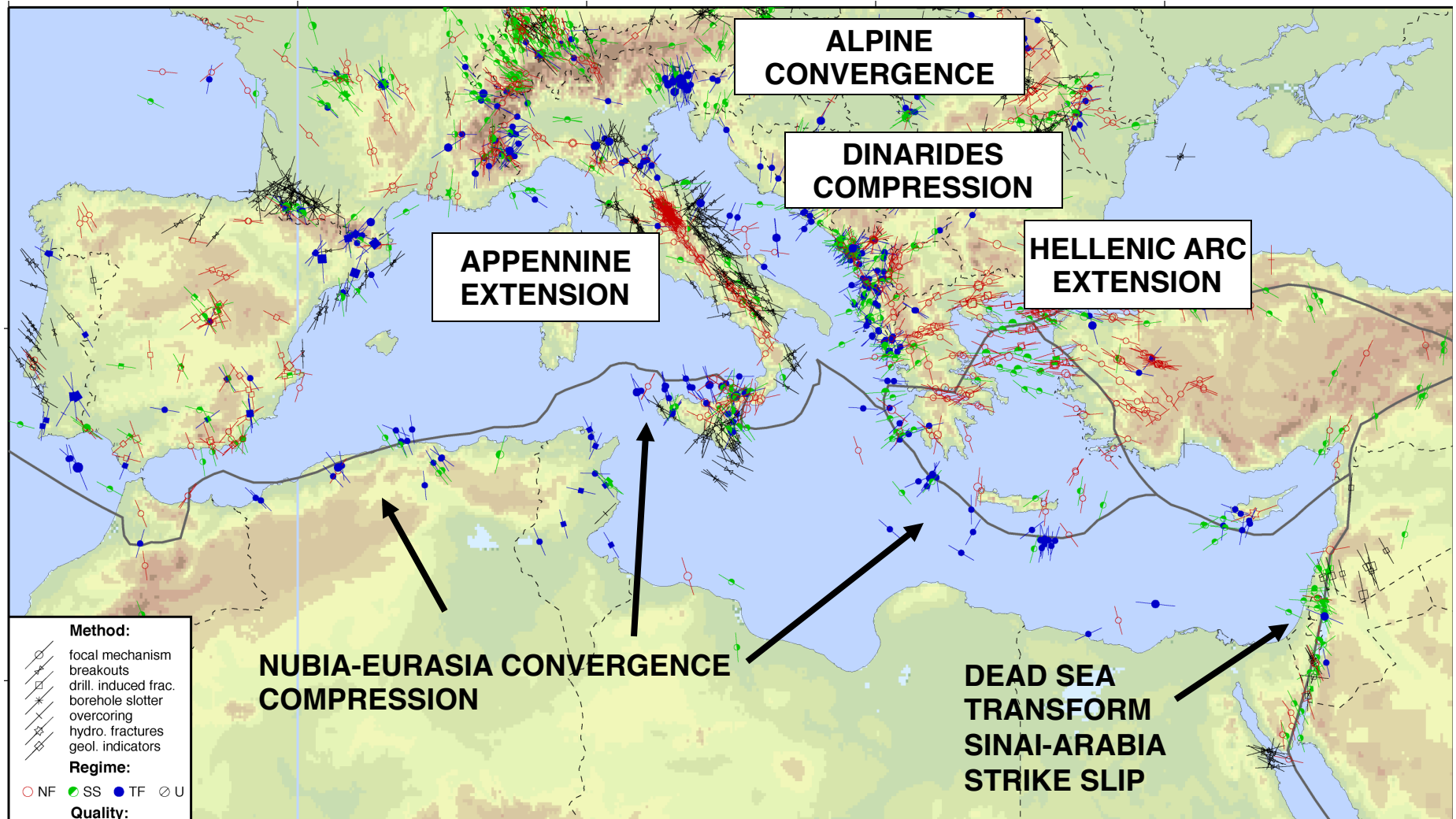
CASCADIA: America's ticking time-bomb



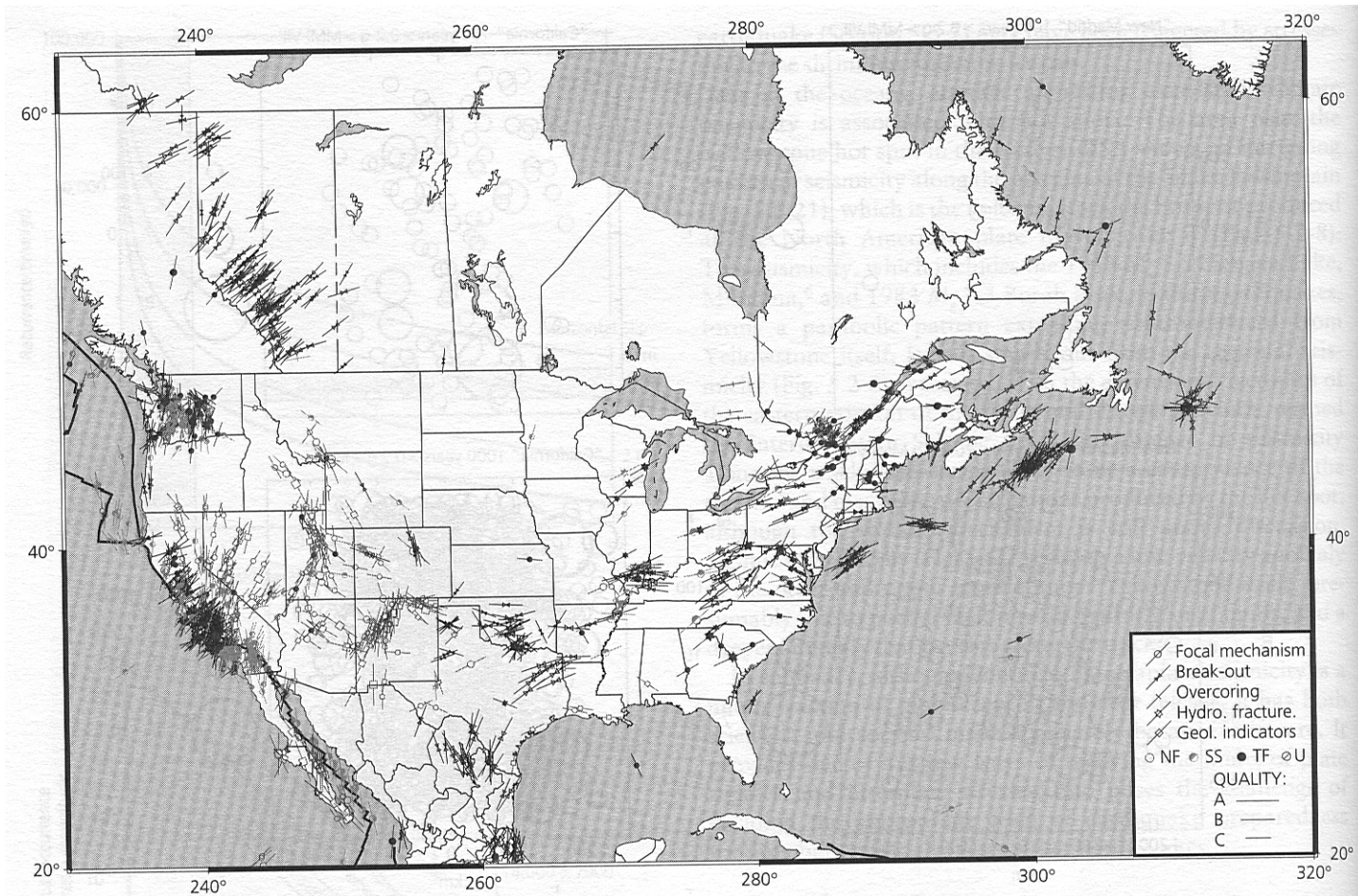
Credit: University of Washington's Center for Environmental Visualization

WORLD STRESS MAP

Combines earthquake mechanisms & other stress indicators



US stress map

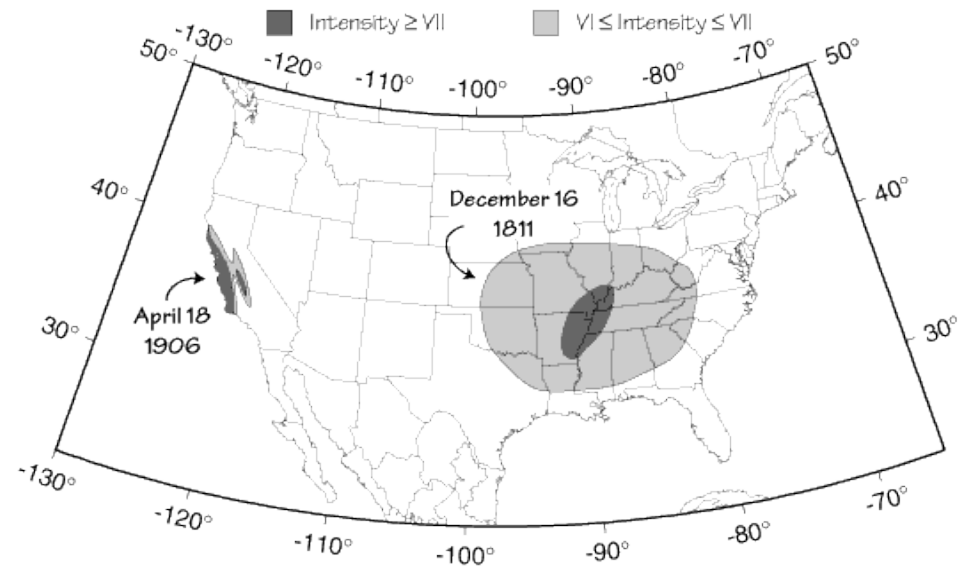
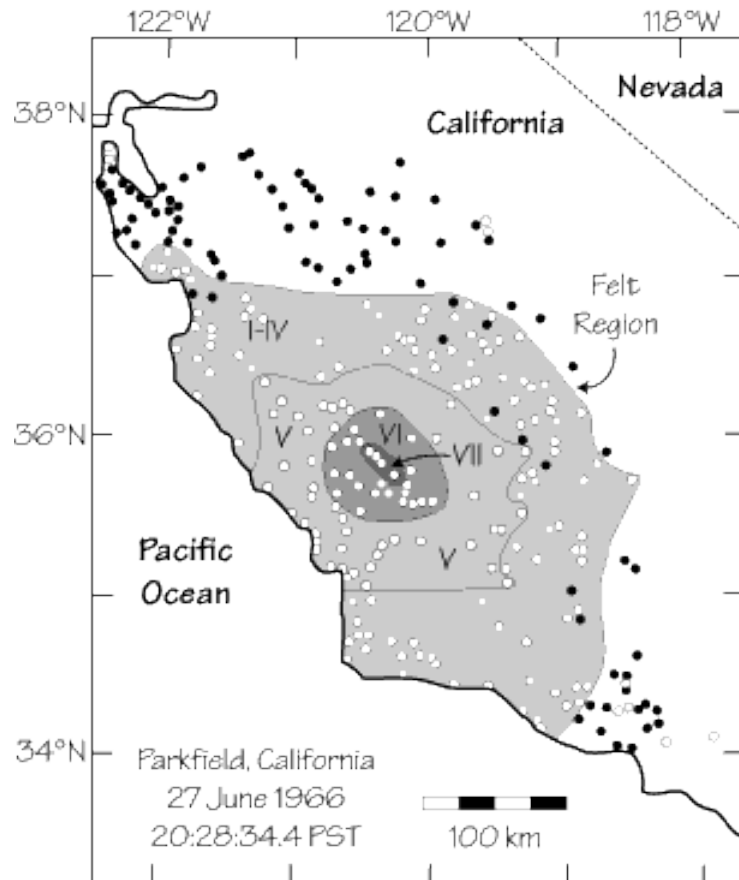


How do we measure the earthquake size?

- Intensity: based on damage

Imperceptible
I II III IV V VI VII VIII IX X Total Destruction
XI XII

Modified Mercalli Scale

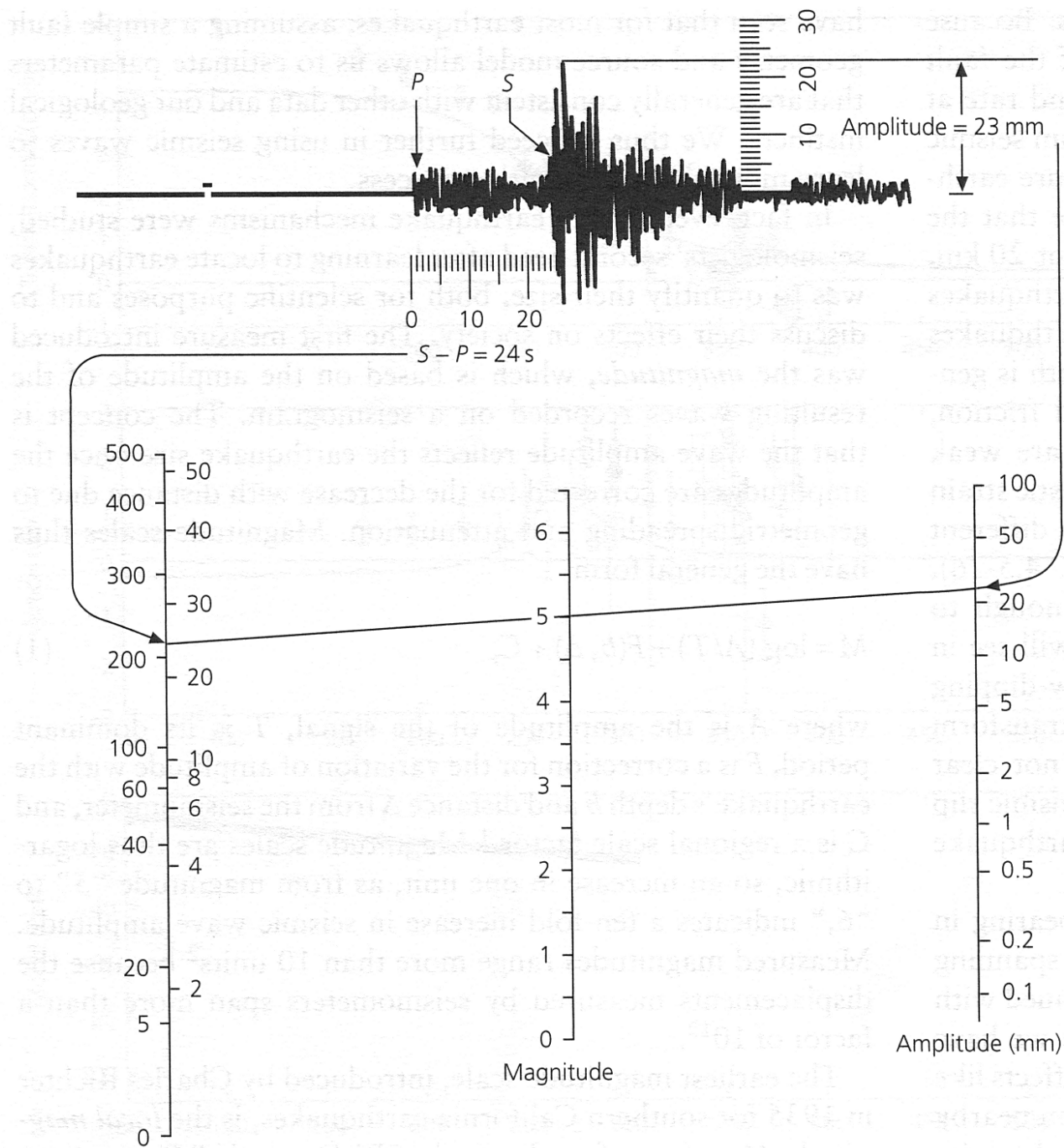


How do we measure the earthquake size?

- Intensity: based on damage
- Magnitude (Richter scale)



L.A. story



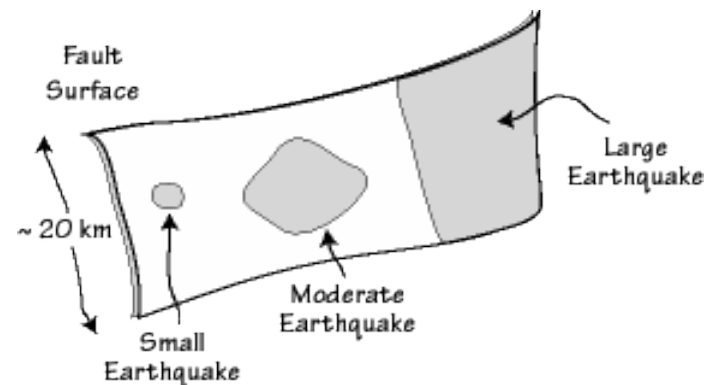
How do we measure the earthquake size?

- Intensity: based on damage
- Magnitude: $M = \log(\text{amplitude}) + a\Delta + b$

Maximum Intensity	I-II	III	IV-V	VI	VII	IX-X	XI	XII
Magnitude	2.0	3.0	4.0	5.0	6.0	7.0	8.0	8.5

How do we measure the earthquake size?

- Intensity: based on damage
- Magnitude: $M = \log(\text{amplitude}) + a\Delta + b$
- Moment: $M_0 = \text{shear modulus} \times \text{rupture area} \times \text{offset}$



1964 ALASKA EARTHQUAKE M_s 8.4 M_w 9.1

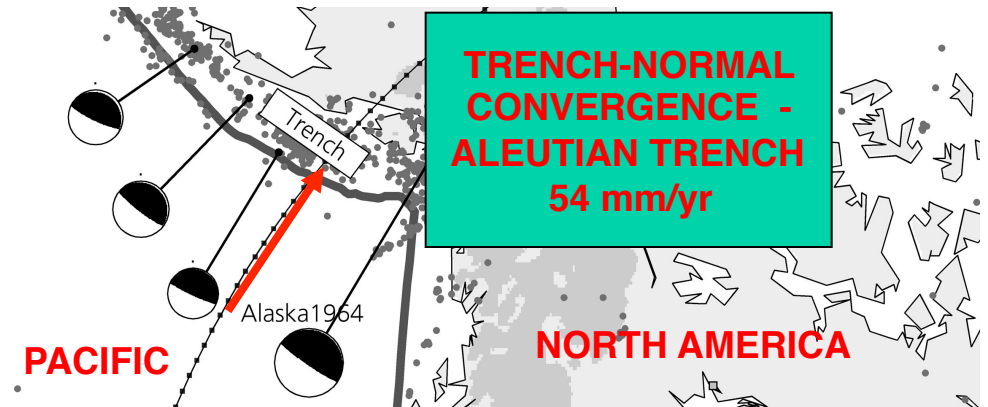
**Pacific subduction
beneath North America**

**~ 7 m of slip on 500x300 km²
of Aleutian Trench**

**Second largest earthquake
recorded until Indonesia
(2004) and Japan (2011)
quakes**

~ 130 deaths

**Catalyzed idea that great
thrust fault earthquakes
result from slip on
subduction zone plate
interface**





1971 M_s 6.6 SAN FERNANDO EARTHQUAKE

1.4 m slip on 20x14 km² fault

Thrust faulting from compression across Los Angeles Basin

Fault had not been previously recognized

65 deaths, in part due to structural failure

Prompted improvements in building code & hazard mapping

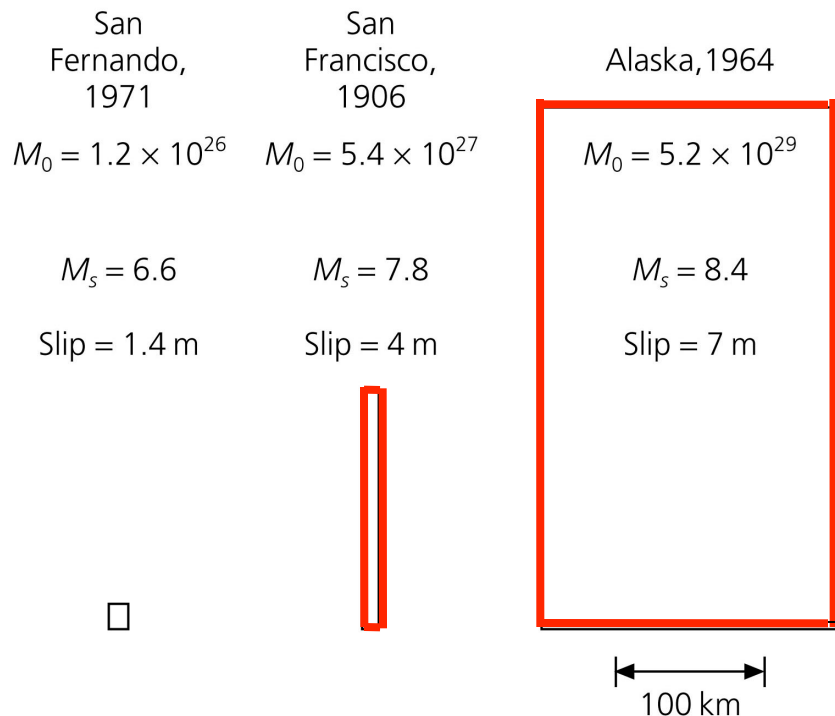


COMPARE EARTHQUAKES USING SEISMIC MOMENT M_0

$$M_0 = \mu \bar{D} S$$

\bar{D} = average slip (dislocation)

S = "average" fault area



Magnitudes, moments (dyn-cm), fault areas, and fault slips for several earthquakes

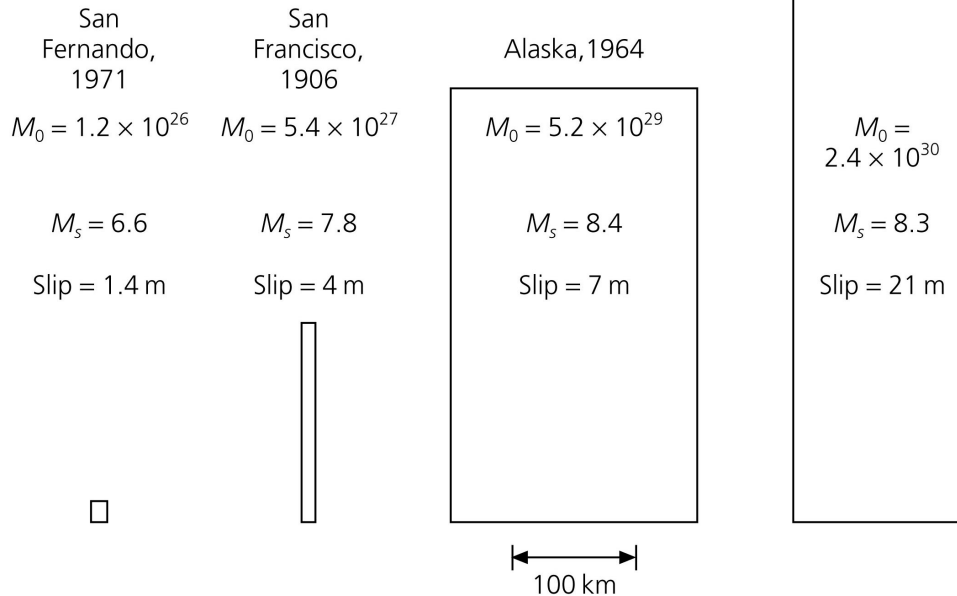
Alaska & San Francisco differ much more than M_s implies

M_0 more useful measure

Units: dyne-cm or N-m

Directly tied to fault physics

Chile, 1960



Moment magnitude M_w

Magnitudes saturate:

No matter how big the earthquake

m_b never exceeds ~ 6.4

M_s never exceeds ~ 8.4

M_w defined from moment so never saturates

Moment magnitude:

$$M_w = \frac{\log M_0}{1.5} - 10.73$$

(with M_0 in dyn-cm)

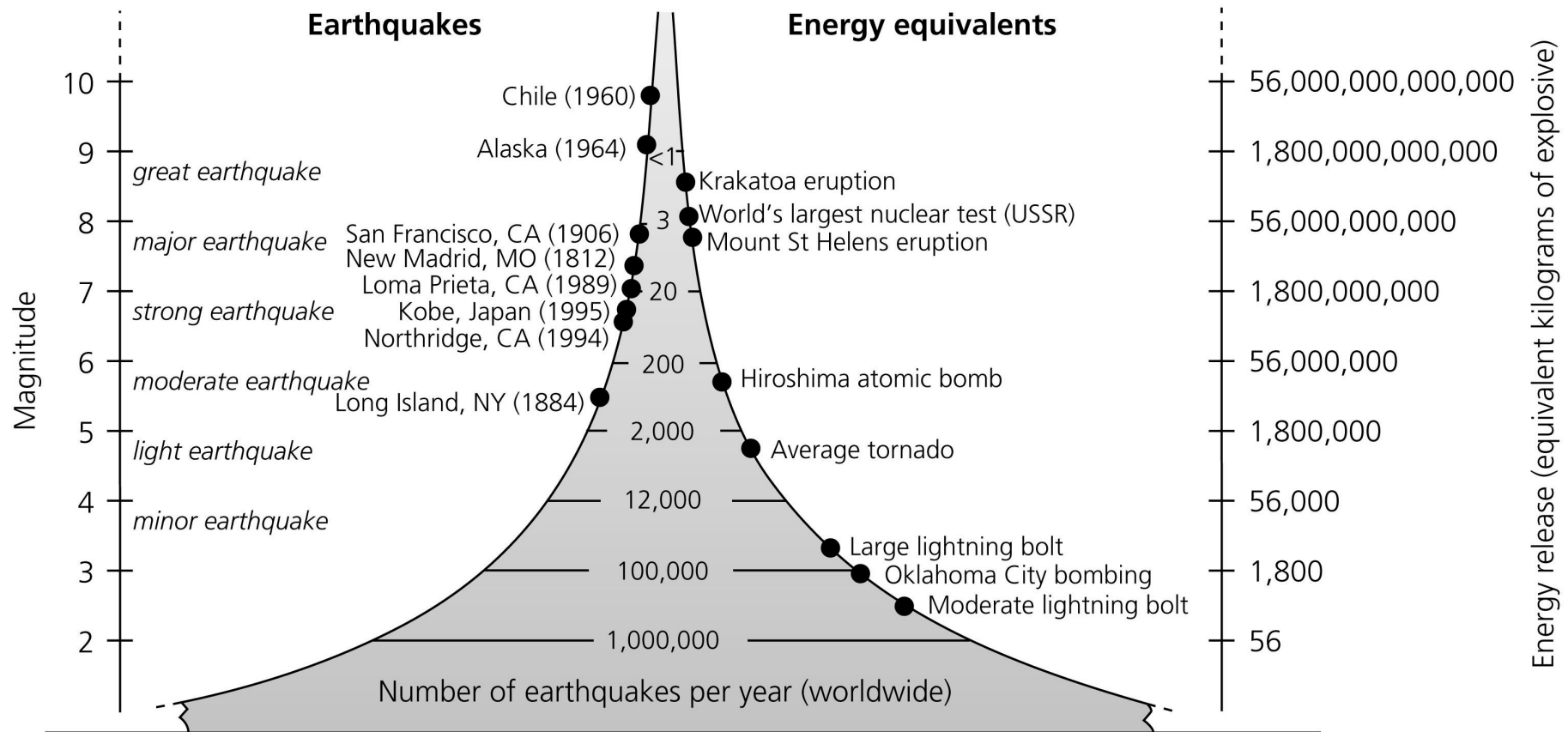
note units!

Earthquake	Body wave magnitude m_b	Surface wave magnitude M_s	Fault area (km ²) length \times width	Average dislocation (m)	Moment (dyn-cm) M_0	Moment magnitude M_w
Truckee, 1966	5.4	5.9	10 \times 10	0.3	8.3×10^{24}	5.8
San Fernando, 1971	6.2	6.6	20 \times 14	1.4	1.2×10^{26}	6.7
Loma Prieta, 1989	6.2	7.1	40 \times 15	1.7	3.0×10^{26}	6.9
San Francisco, 1906		8.2	320 \times 15	4	6.0×10^{27}	7.8
Alaska, 1964	6.2	8.4	500 \times 300	7	5.2×10^{29}	9.1
Chile, 1960		8.3	800 \times 200	21	2.4×10^{30}	9.5

How do we measure the earthquake size?

- Intensity: based on damage
- Magnitude: $M = \log(\text{amplitude}) + a\Delta + b$
- Moment: $M_0 = \text{shear modulus} \times \text{rupture area} \times \text{offset}$
- Energy: $\log(E) = 1.5 M + 5.2$

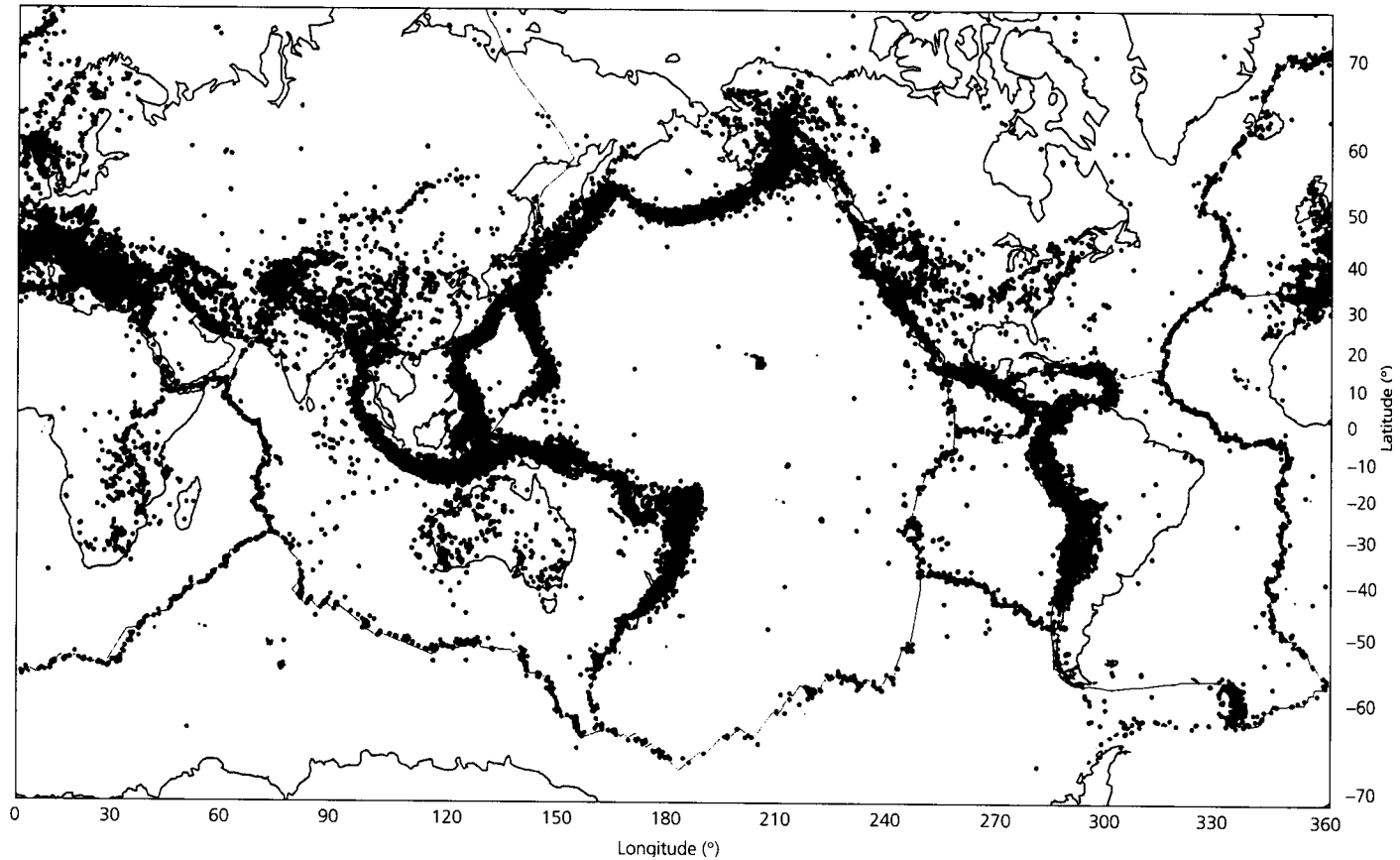
Comparison of frequency, magnitude, and energy release.



Earthquakes of a given magnitude are ~10 times less frequent than those one magnitude smaller. An M7 earthquake occurs approximately monthly, and an earthquake of M> 6 about every three days. Hence although earthquake predictor I. Browning claimed to have predicted the 1989 Loma Prieta earthquake, he said that near a date there would be an M6 earthquake somewhere, a prediction virtually guaranteed to be true.

Magnitude is proportional to the logarithm of the energy released, so most energy released seismically is in the largest earthquakes. An M 8.5 event releases more energy than all other earthquakes in a year combined. Hence the hazard from earthquakes is due primarily to large (typically magnitude > 6.5) earthquakes.

Where is most of seismic energy released?



Shallow Earthquake Moment Release
1900-1989

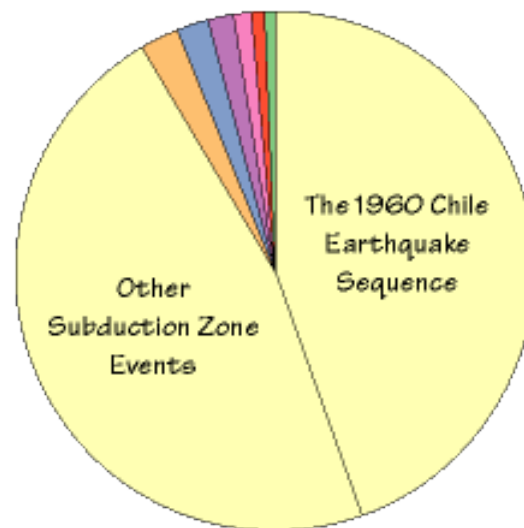
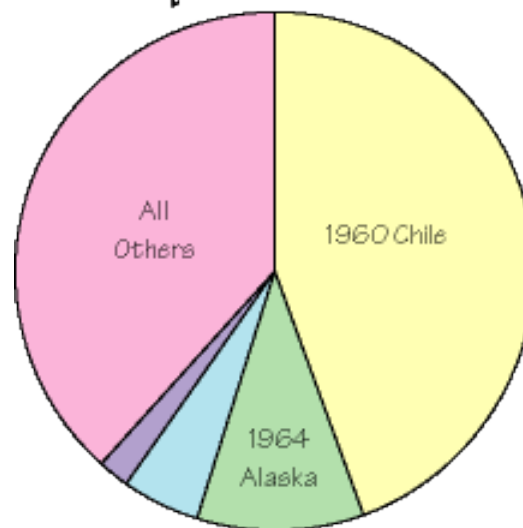


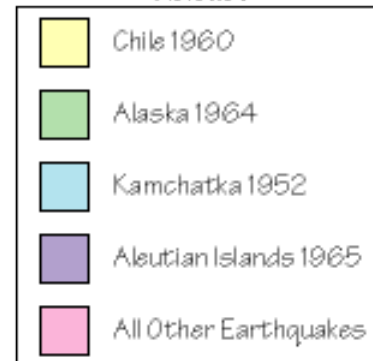
Plate Boundary Type



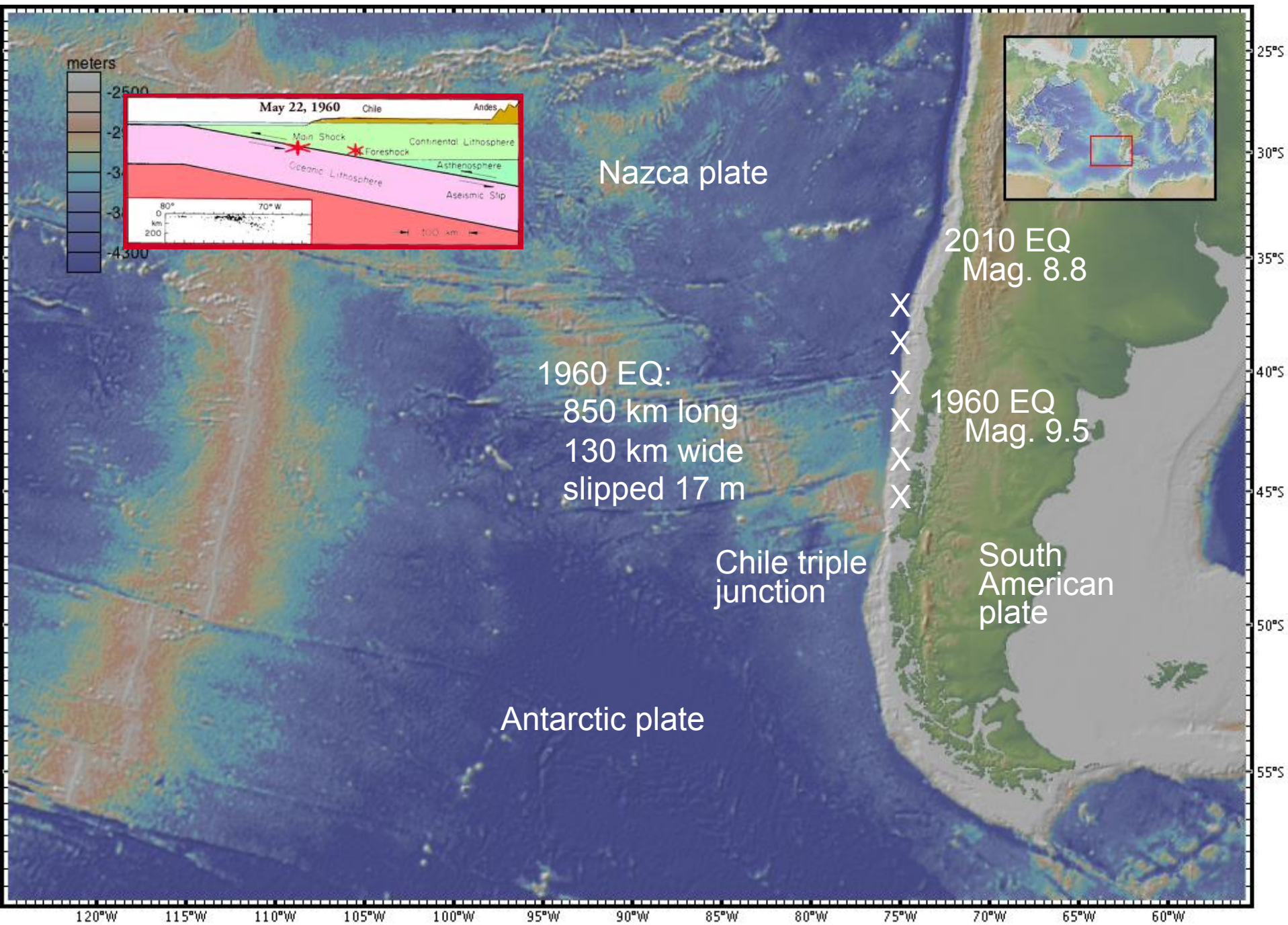
Giant Shallow Earthquakes 1900-1998

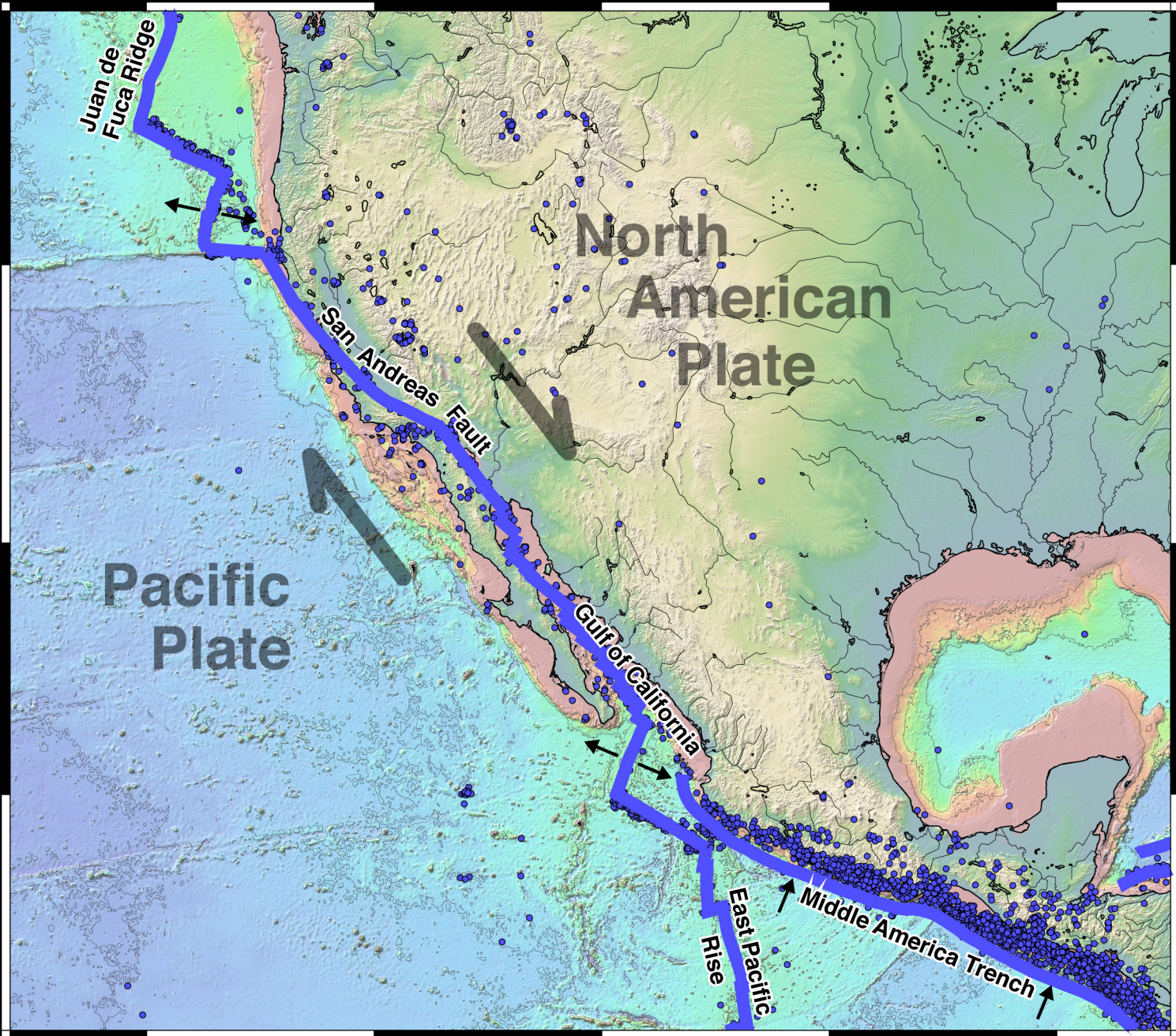


Earthquake Moment Release



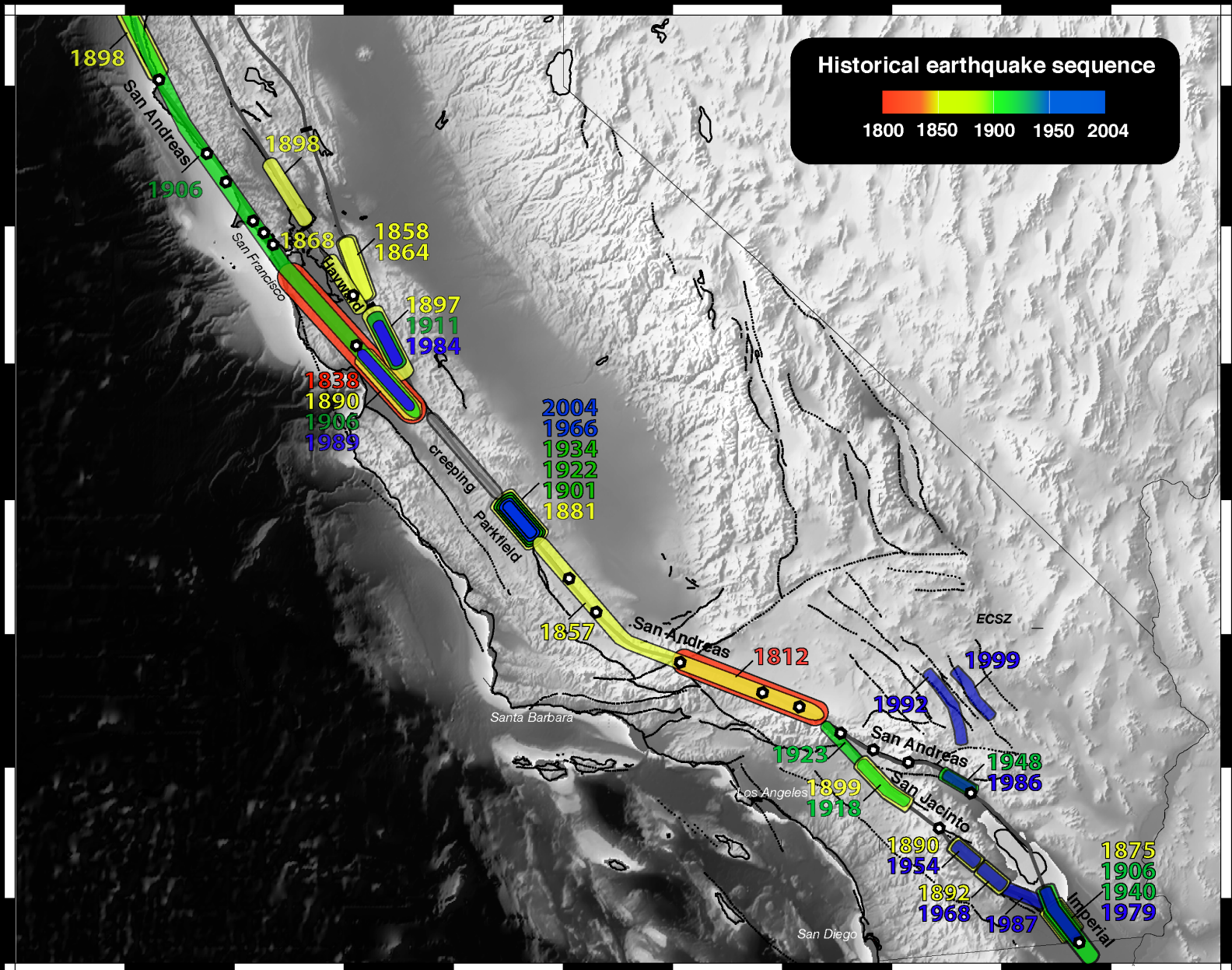
Source: Pacheco and Bykes, 1992





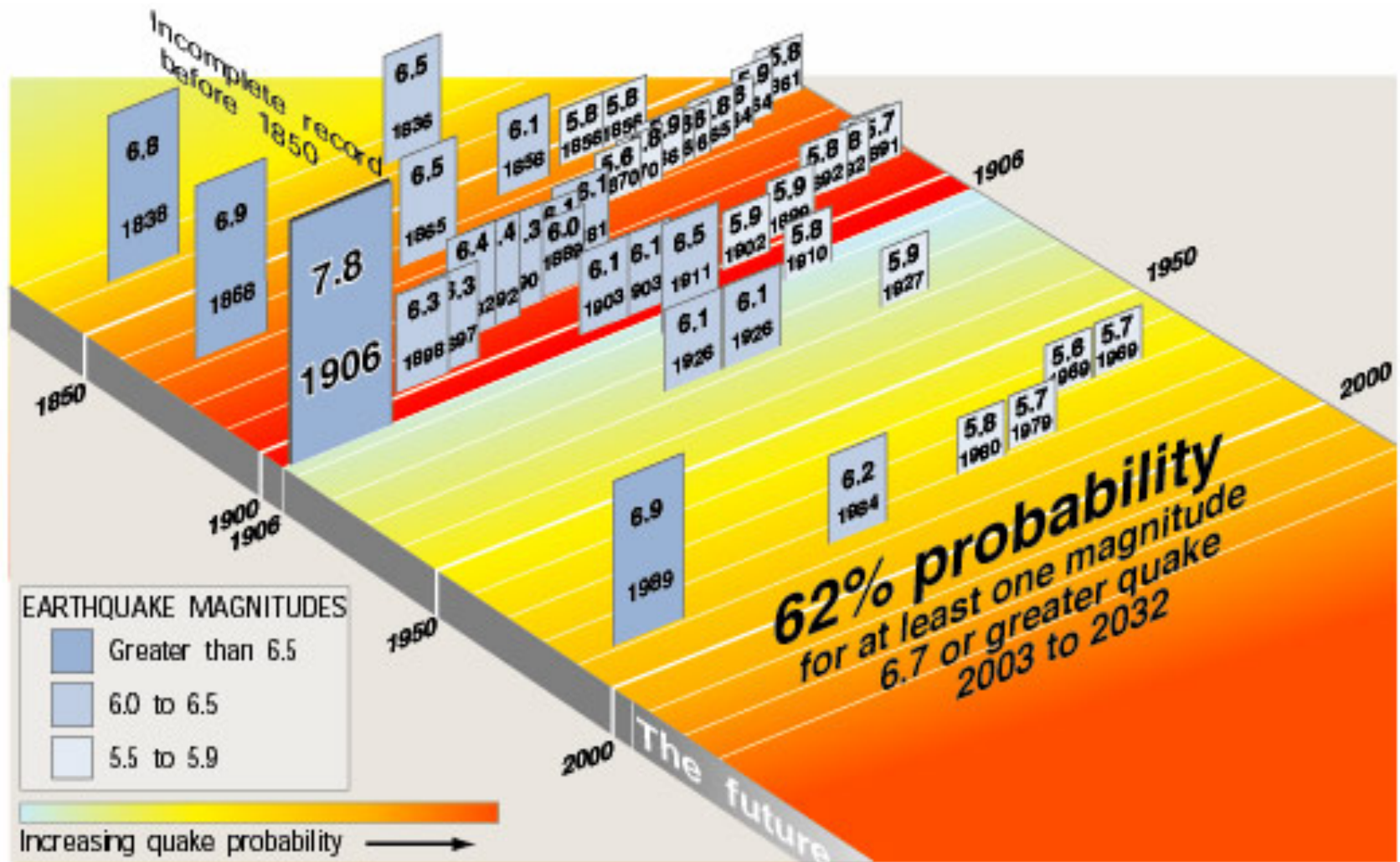


Pre-historical Earthquakes (1000-1800)

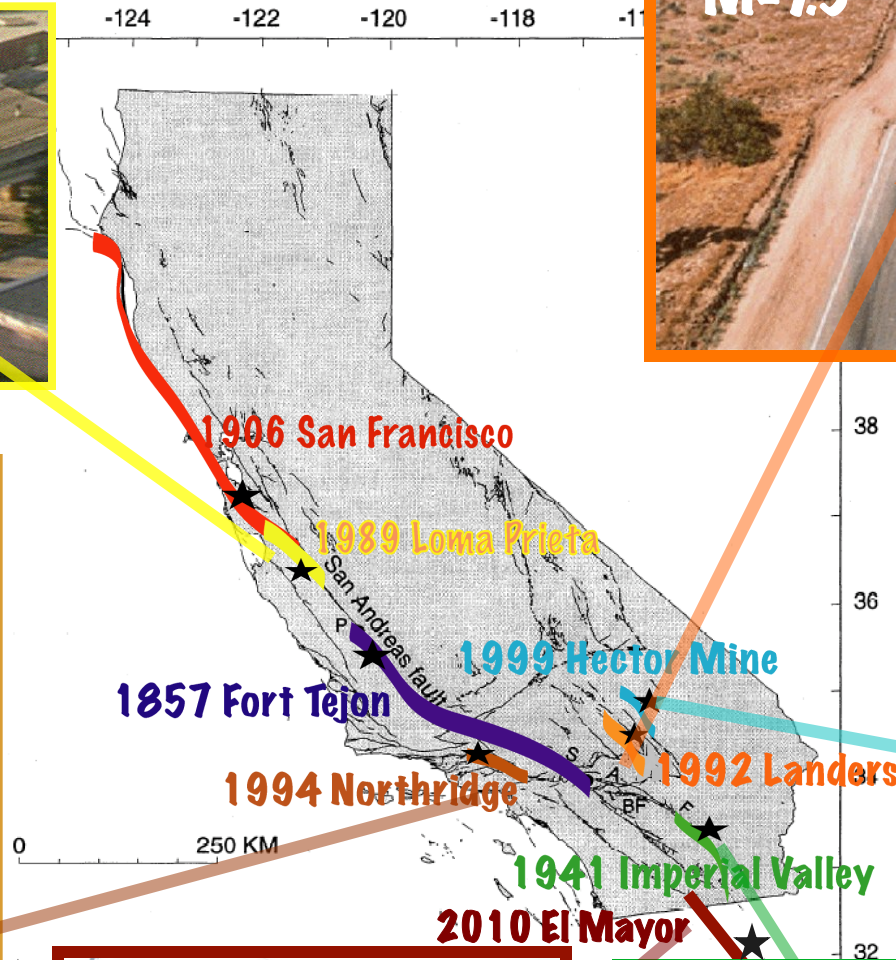
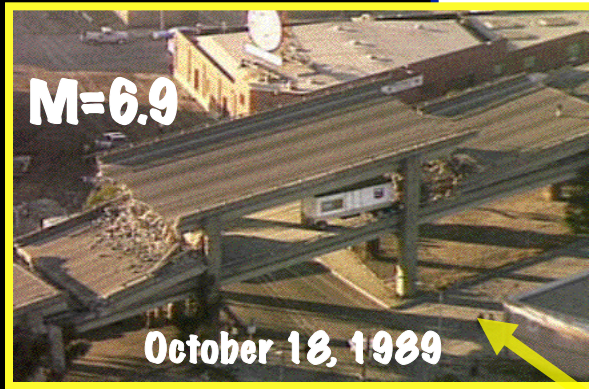


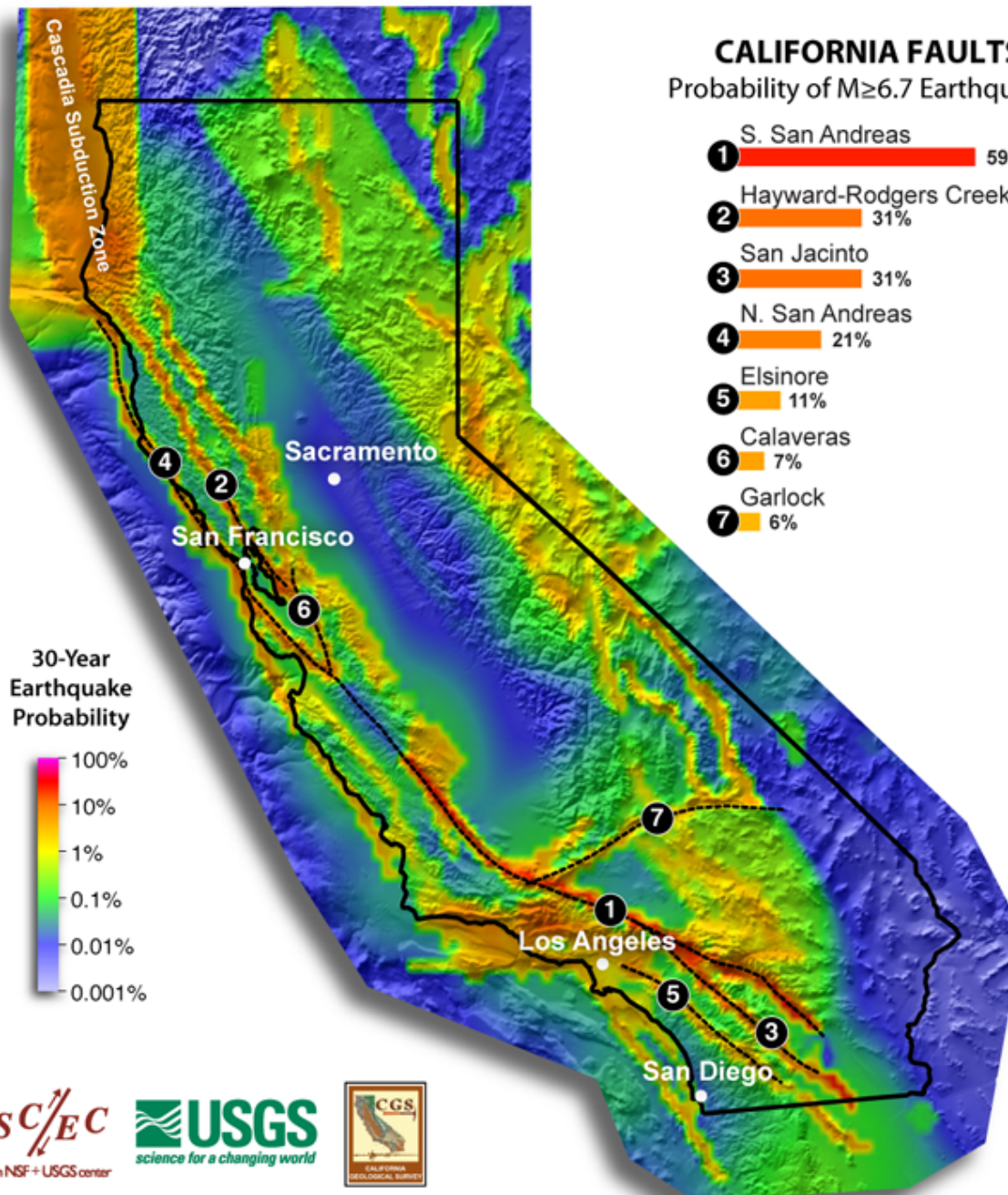


Significant earthquakes in the San Francisco Bay Area: before and after 1906

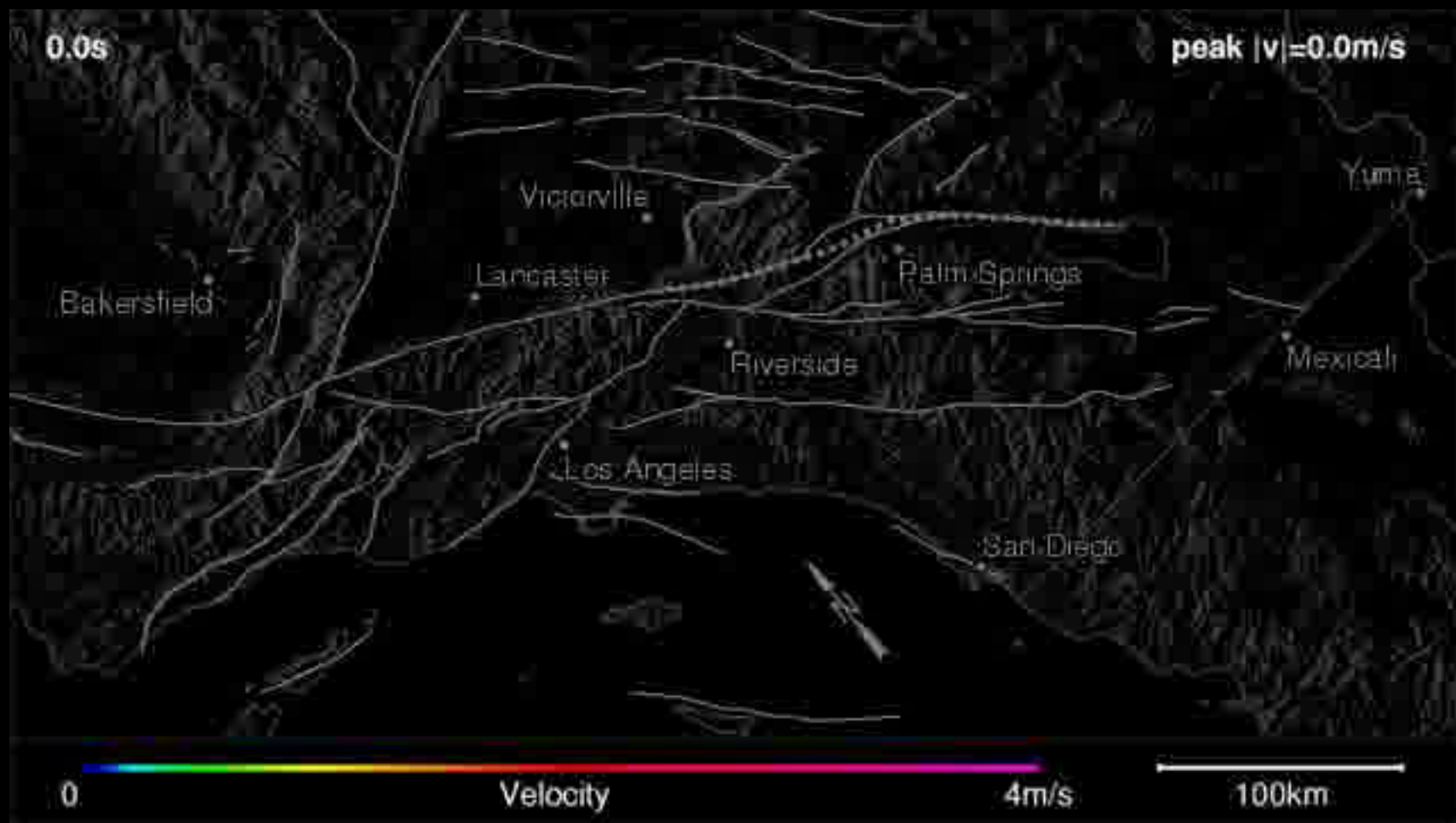


Recent large EQs in CA

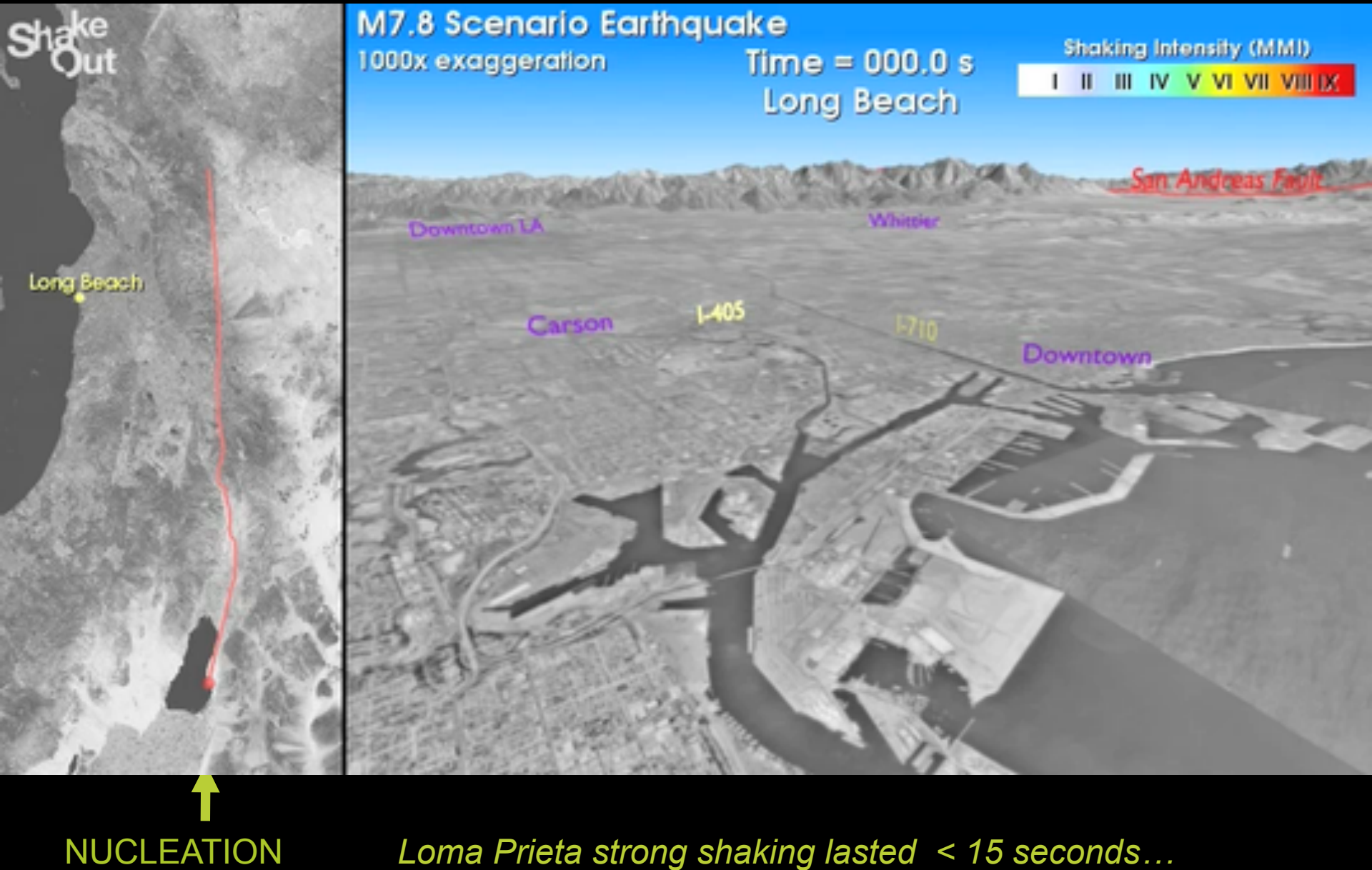








Simulation of Ground Motion



Onset and Duration of Shaking

Location	Seconds after start of earthquake that strong shaking begins at this location	Seconds after start of earthquake that strong shaking ends at this location	Duration of very strong shaking
Palm Springs	25	60	35 sec
San Bernardino	45	75	30 sec
Los Angeles (downtown)	70	125	55 sec
Orange County	70	105	35 sec
Santa Monica	85	150	65 sec
Palmdale	75	90	15 sec
Ventura	105	160	55 sec

ShakeOut Scenario “Disaster Equation”

Widespread Strong Ground Shaking
+ Shaking of Long Duration =

300,000 buildings significantly damaged

Widespread infrastructure damage

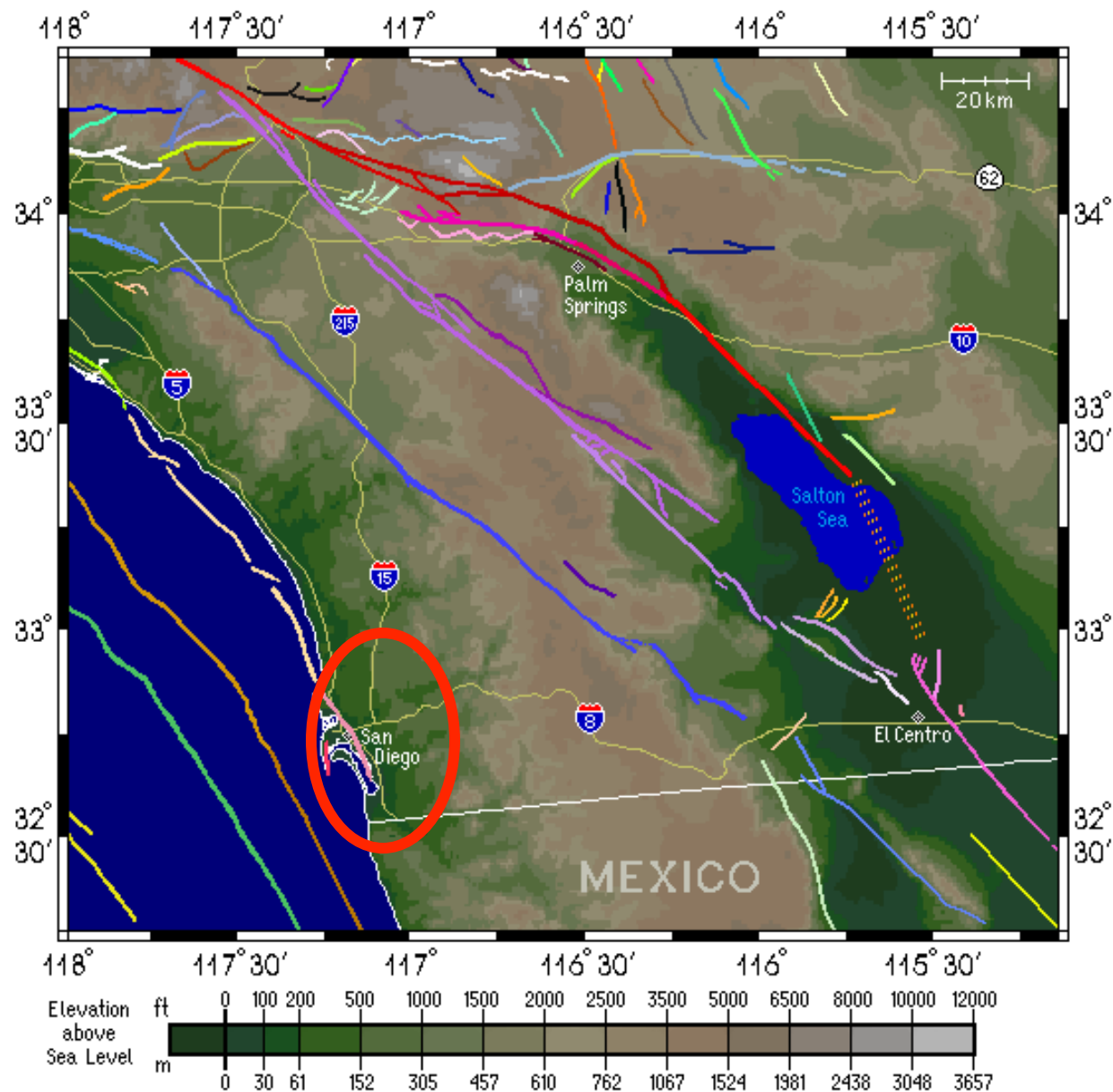
\$213 billion damages

270,000 displaced persons

50,000 injuries

1,800 deaths

Local hazards



Rose Canyon - Newport Inglewood fault

Slip rate: ~ 1 mm/yr

Recurrence interval: unknown

Potential for M 6-7 events

“near field”:

high-frequency,

high g

shaking



The Great Southern California ShakeOut

- Every Fall: ShakeOut drill
- 7th year
- A day of special events to inspire southern Californians to get ready for big earthquakes
 - millions of participants: schools, families, community groups, business, etc.
-
- www.shakeout.org

Design by Art Center College of Design

Shake Out. Don't Freak Out.



DROP! COVER! HOLD ON!

October 21, 10:21 a.m.

The Great California **Shake Out**™

Register at www.ShakeOut.org

© 2010 ECA