Seismic Scattering in the Deep Earth

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Mantle mixing calculations



Heterogeneity is likely at all scales

Davies (2002)

Xie & Tackley (2004)

Mantle tomography constraints



Global models dominated by long-wavelength structure (e.g., *Su & Dziewonski*, 1991)



1225 km depth slice (SIO 2004 model)

- But mid-mantle has whiter spectrum
- Resolution limited to 500 to 1000 km

Scattering to the rescue....



- High-frequency waves are scattered by small-scale structures
- Origin of coda in short-period seismograms
- Modeled with random heterogeneity

But strong near-surface scattering complicates study of deep mantle



Deep mantle scattering is hard to study because it is masked by much stronger scattering at shallow depths.



Deep Earth scattering observations

- *1. P* coda
- 2. Pdiff coda
- 3. PP precursors
- 4. P'P' precursors
- 5. PKP precursors
- 6. PKKP precursors
- 7. *PKKPx*
- 8. PKiKP coda



Uniquely valuable *PKP* precursors

 Core *P* velocity drop bends rays so that scattered waves can arrive *before* direct phases ÷.

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TUC 134.7° MMMMMMMMM	M	vww		WMM	MMMM
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PKP Precursor Examples

Higher amplitudes at longer range, where scattering angles are less Stacking complications at short periods...



Stacking incoherent waves



~1 Hz seismograms are incoherent

Compute envelope functions

Stack envelope functions, correcting for energy in pre-event noise

Average PKP Precursor Wavefield



PKP(DF) Precursors

Precursor onset time agrees with CMB as base of scattering (no outer core scattering)



Time





Predicted *PKP* precursor envelopes for scattering at different depths

Stacks suggest whole mantle scattering







PKP Precursor Interpretation



- ~0.5 to 1% RMS velocity perturbations at ~10 km scale length
- Recent analyses show scattering extends at least 1000 km above CMB (Hedlin et al., 1997; Cormier, 1999; Margerin & Nolet, 2003).
- Early studies put scattering near CMB (e.g., *Cleary & Haddon*, 1972)

P_{diff} coda provides more evidence for mid-mantle scattering

coda





Earle & Shearer (2001)

Time before PP (s) -250 - 300 -350 Range (degrees)

Good fit to data stack obtained with 1% RMS velocity heterogeneity throughout the mantle *P*_{diff}

ScS coda analysis supports lower mantle scattering





Lee et al. (2003) used radiative transfer modeling

Largest scattered signal in highfrequency wavefield is *P* coda



coda

Whole Earth Scattering: A Challenging Modeling Problem



Synthetics should include:

- Strong scattering (multiple)
- Weak scattering (single, Born)
- P and S waves
- Random perturbation models
- Reflection/transmission coef.
- Geometrical spreading
- Intrinsic attenuation
- Energy conservation

Monte Carlo seismic "photon" method

Spray particles from source



Randomly scatter using probabilities computed from random media theories.

> Gusev and Abubakirov (1987) Hoshiba (1991, 1994, 1997) Yoshimoto (2000) Margerin et al. (2000)

(Long used in physics, related to radiative transfer theory)

Scattering in random media



- These models have just two parameters: correlation distance and RMS velocity perturbation
- Born theory gives scattering power per volume and power at different scattering angles
- Easy to convert to particle probability, mean free path

Energy partitioning at interfaces and scattering regions is handled as probability for change in photon path.



Four photons go through (on average)







Advantages of Method

- Energy conservation is maintained
- No need to specify all possible ray paths
- Computes *complete* wavefield
- Includes both intrinsic and scattering Q
- Includes multiple scattering
- Handles both volume and interface scattering

Limitations

- Ray theory doesn't get diffracted waves
- Polarity/phase information is lost



IASP91 travel time curves

- Assume equal partitioning between reflected & converted waves at surface, CMB and ICB
- No attenuation
- Automatically generates all possible travel time curves for P and S
- Actual amplitudes for most are too small to see in data



from Shearer & Earle (2004)

Synthetic fit to P coda requires some lower mantle scattering





Surface source, 3-layer scattering model

- 4% velocity variation above 200 km
- 3% variation from 200 to 600 km
- 0.5% heterogeneity below 600 km
- Higher *Q* than *Warren & Shearer* (2001)
- 4-km scale length in upper mantle 8-km scale length in lower mantle

from Shearer & Earle (2004)

Same model fits both shallow and deep earthquake results



New results from LASA array: *P'P'* scattering at short distances



from Earle & Shearer (2007)



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from *Earle & Shearer* (2007)





from Earle & Shearer (2007)

Ray geometry for *P'-d-P'* scattering







Onset time for *P'P'* scattering from 400 km

Modeling P'dP' scattering at short distances may provide best insight yet into depth dependence of high-frequency scattering in the mantle



Conclusions

- Many different seismic studies indicate small-scale (~10 km) random heterogeneity in deep mantle
- RMS amplitude is still an issue (*Margerin & Nolet* get much smaller number than *Hedlin et al.*)
- Needed: Analysis at different frequency bands to constrain power spectrum of heterogeneity over 1 to 500 km band
- Implications for geochemistry and convection modeling should be explored
- Monte Carlo code should be useful for modeling *lots* of additional scattered phases, including newly discovered *P'dP'* scattering



